



Crustal Thickness Estimatives and Vp/Vs Ratio Using Receiver Functions.

Diogo Farrapo Albuquerque, César Garcia Pavão, Rafael Toscani Gomes da Silveira, Iago Guilherme dos Santos and George Sand França (IG-SIS-UnB).

Copyright 2011, SBGf - Sociedade Brasileira de Geofísica

This paper was prepared for presentation during the 12th International Congress of the Brazilian Geophysical Society held in Rio de Janeiro, Brazil, August 15-18, 2011.

Contents of this paper were reviewed by the Technical Committee of the 12th International Congress of the Brazilian Geophysical Society and do not necessarily represent any position of the SBGf, its officers or members. Electronic reproduction or storage of any part of this paper for commercial purposes without the written consent of the Brazilian Geophysical Society is prohibited.

Abstract

Receiver Functions calculated by frequency domain deconvolutions have been widely used to estimate crustal thickness and Vp/Vs ratio. The aim of this research is to obtain such estimatives using seismographic data recorded by seven stations belonging to Seismological Observatory: BRA7 (Brasília-DF), CAN3 (Cana Brava-GO), FOR1 (Fortaleza-CE), JAN7 (Itacarambi-MG), PDRB (Porto dos Gaúchos-MT), SFA1 (Serra do Facão-GO) e TUCA (Tucuruí-PA). In addition, all the results complement and corroborate the crustal thickness values already obtained for several tectonic provinces and it helps us to understand the processes involved in the tectonic formation of Brazilian territory.

Introduction

Receiver Function (RF) is a geophysical technique used mainly to study the Earth structure employing teleseismic events located at epicentral distances ranging from 30° to 90°. This technique is very important for determining the depth of Moho discontinuity beneath a seismographic station.

When the P-wave reaches a discontinuity, part of its energy is converted into S (Ps-wave) and also in multiple reflections (Ppps and Ppsps+Pps). The incidence angle of the seismic waves in the mantle-crust interface are typically smaller than 40°, which would require a record of the P wave dominant in the vertical, and that converted to S (Ps) predominant in the radial component. The RFs are obtained by the deconvolution of the radial component from the vertical. The deconvolution process isolates the receptor response and then the result contains the signature of the converted phases and multiple reflections (Langston, 1979; Owens, 1984; Ammon, 1991).

Method

Teleseismic events located at epicentral distances ranging from 30° to 90° (Figure 1) and magnitudes equal or greater than 4.5 m_b were selected. All records were obtained by seismographic stations belonging to Seismological Observatory, Brasília University (SIS-UnB). The data were converted from GCF (Güralp Systems format) to SAC (Seismic Analysis Code; Goldstein &

Snoke, 2005). The teleseismic waveforms were selected using the program *getevts* (An, 2004), which searches for information about earthquakes in the IRIS (Incorporated Research Institutions for Seismology) network. After choosing the waveforms, it was made a visual inspection marking the P-wave time in the best events. For this analysis, it was used the SAC software (Goldstein & Snoke, 2005). The deconvolution in frequency domain was performed using the program *pwaveqn* (Ammon, 1991). All the RFs were plotted according to their azimuths and epicentral distances of the events. The stacking was performed by the program *HK-Stacking* (Zhu & Kanamori, 2000). This technique uses the relative moveout correction of secondary arrivals generated at a seismic discontinuity beneath the station to infer the depth and the Vp/Vs ratio above the discontinuity (França, 2003).

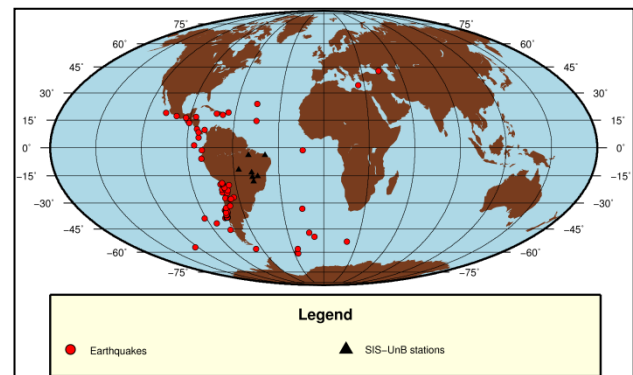


Figure 1 – Teleseismic events and location of SIS-UnB stations.

Results

The crustal thickness estimatives and Vp/Vs ratios are shown in the Table 1.

Table 1 – Crustal thickness and Vp/Vs ratios.

Station	Depth (Km)	Vp (Km/s)	Vp/Vs
BRA7	41.7 ± 0.7	6.3	1.69 ± 0.02
CAN3	40.0 ± 1.1	6.3	1.64 ± 0.02
FOR1	32.4 ± 0.7	6.3	1.75 ± 0.02
JAN7	40.3 ± 1.0	6.4	1.70 ± 0.02
PDRB	37.6 ± 1.5	6.4	1.68 ± 0.05
SFA1	37.1 ± 2.7	6.4	1.69 ± 0.07
TUCA	30.6 ± 4.5	6.4	1.79 ± 0.10

The figures 2 to 5 show the best Receiver Functions obtained using the stations BRA7, CAN3, FOR1 and JAN7. The direct P wave and the refracted P wave (Ps) are indicated by the red boxes.

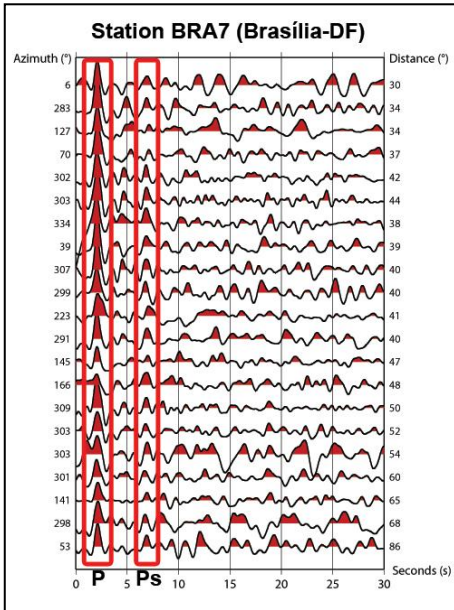


Figure 2 – Receiver Function result for BRA7 station plotted according the azimuths and epicentral distances of the events.

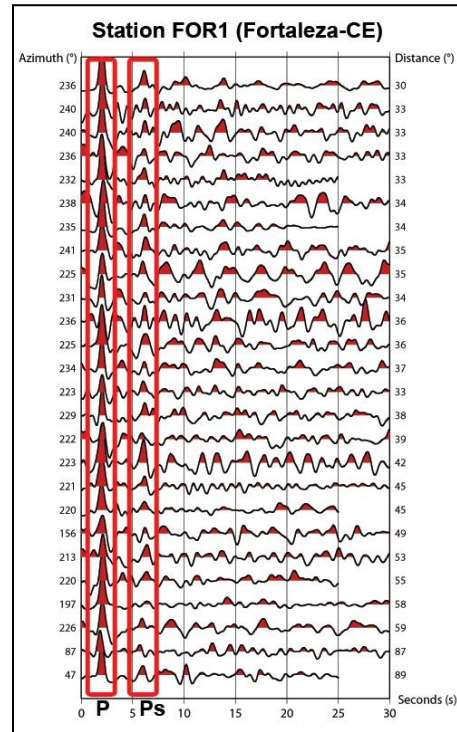


Figure 4 - Receiver Function result for FOR1 station plotted according the azimuths and epicentral distances of the events.

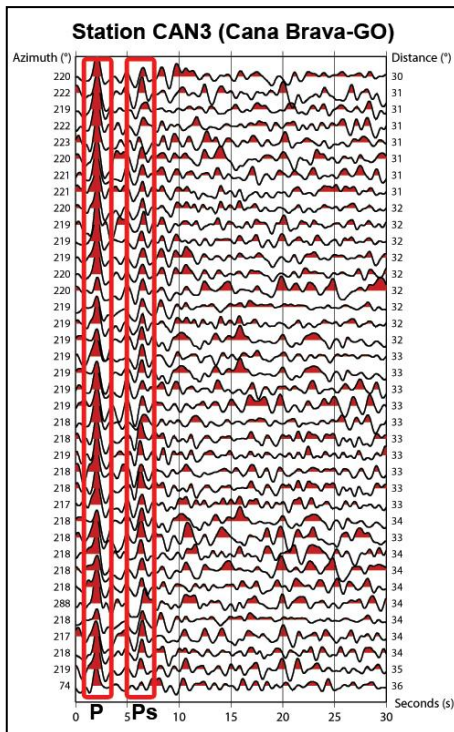


Figure 3 - Receiver Function result for CAN3 station plotted according the azimuths and epicentral distances of the events.

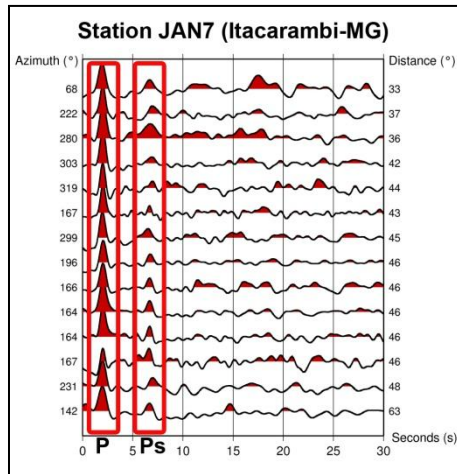


Figure 5 - Receiver Function result for JAN7 station plotted according the azimuths and epicentral distances of the events.

The figures 6 to 9 show the HK-Stacking results obtained using the stations BRA7, CAN3, FOR1 and JAN7. The crustal thickness values are 41.7 ± 0.7 , 40.0 ± 1.1 , 32.4 ± 0.7 , 40.3 ± 1.0 , respectively.

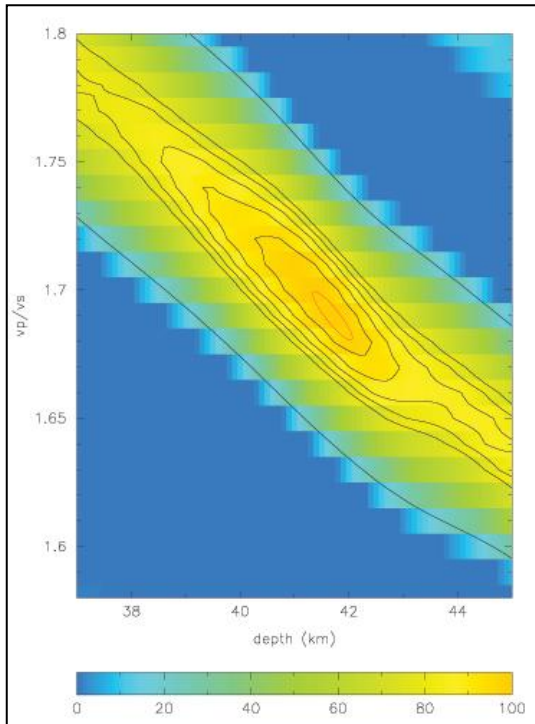


Figure 6 – HK-Stacking result for BRA7. The red ellipse indicates the variation in the value of crustal thickness and V_p/V_s ratio.

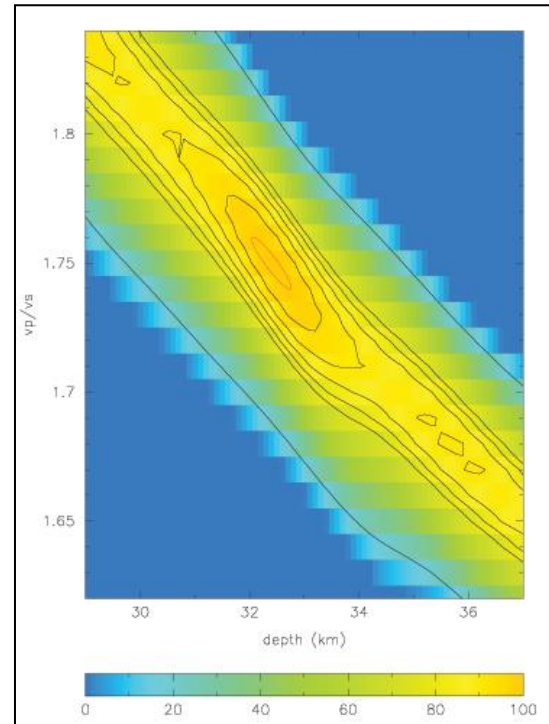


Figure 8 – HK-Stacking result for FOR1. The red ellipse indicates the variation in the value of crustal thickness and V_p/V_s ratio.

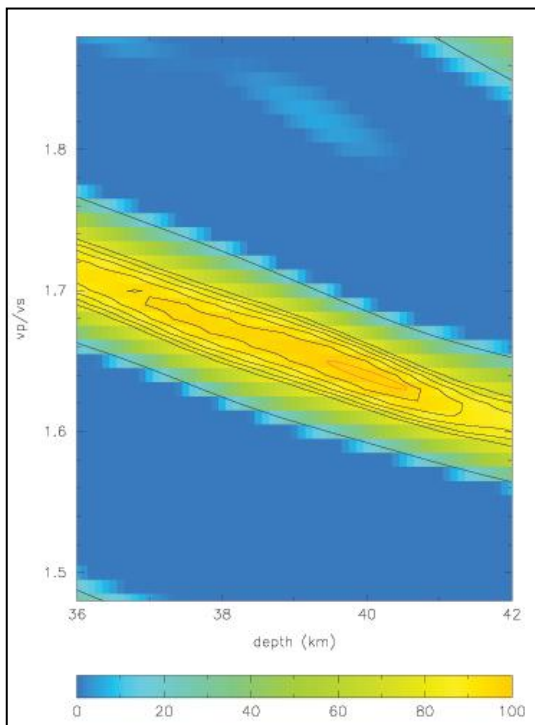


Figure 7 – HK-Stacking result for CAN3. The red ellipse indicates the variation in the value of crustal thickness and V_p/V_s ratio.

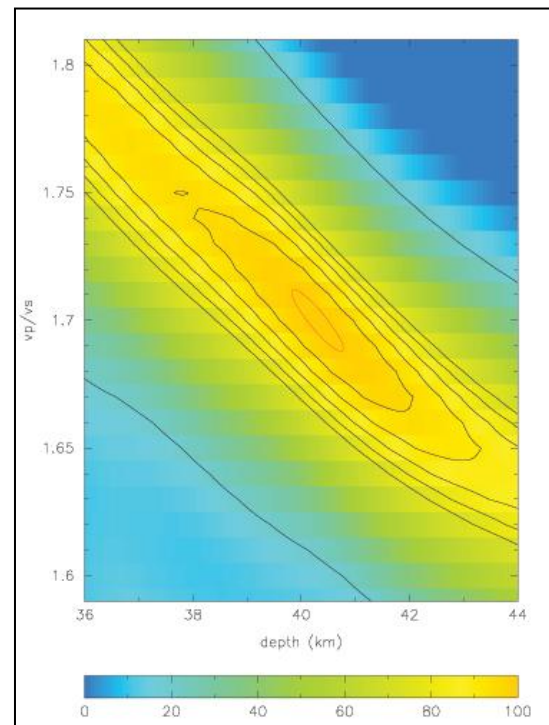


Figure 9 – HK-Stacking result for JAN7. The red ellipse indicates the variation in the value of crustal thickness and V_p/V_s ratio.

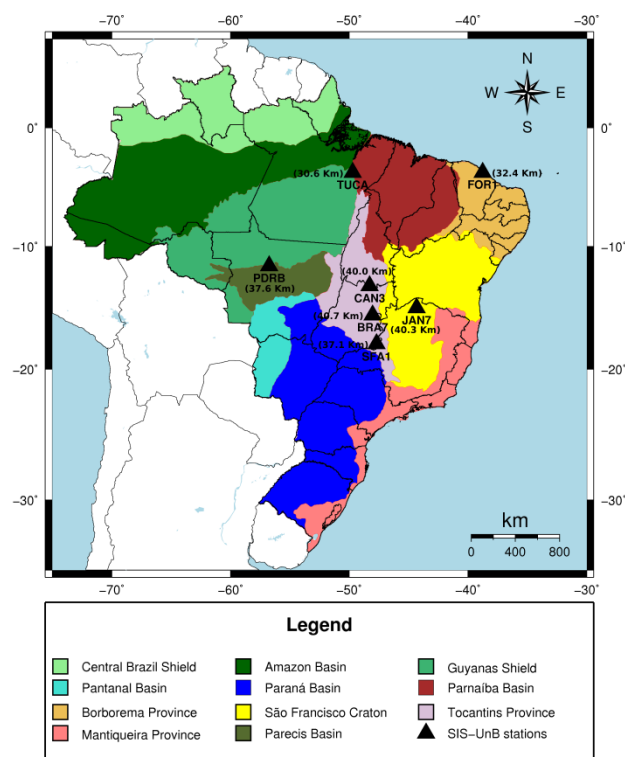
Discussion and conclusions

The crustal thickness and Vp/Vs results for the stations BRA7, CAN3, FOR1 and JAN7 are consistent with expected values in comparison to other previous studies which had been done using stations near or on the same tectonic province (Bianchi, 2008). The best results were obtained for stations placed in cratonic provinces (e.g. JAN7 and FOR1).

For the stations PDRB, SFA1 and TUCA the values are not in accordance with previous studies. Some possible causes are:

- Poor quality of recorded data (high background noise);
- Small amount of teleseismic waveforms;
- Gaps in the data period.

The Figure 10 shows the tectonic provinces, the crustal thickness and Vp/Vs ratio results and the stations belonging to SIS-UnB.



Acknowledgments

This research was supported by Seismological Observatory of Brasília (SIS-UnB) and CNPq.

References

- Ammon, C.J., 1991. The isolation effects from teleseismic P-waveforms. *Bull. Seism. Soc. Am.* 81, 2504–251.
- Ammon, C. J., G. E. Randall & G. Zandt, 1990. On the nonuniqueness of receiver functions inversions, *J. Geophys. Res.*, 95, 15303-15318.

An, M., and M. Assumpção (2004), Crust and upper mantle structure in intracratonic Paraná basin from surface wave dispersion using genetic algorithm, submitted to *J. South Am. Earth Sci.*

Assumpção, M., An, M., Bianchi, M., França, G.S.L., Rocha, M., Barbosa, J.R., Berrocal, J., 2004. Seismic studies of the Brasília Fold Belt as the western border of the São Francisco Craton, Central Brazil. *Tectonophysics*, 388: 173-185.

Bianchi, M., 2008. Variação da estrutura da crosta, litosfera e manto para a plataforma Sul Americana através de funções do receptor para ondas P e S. Tese de Doutorado, Instituto de Astronomia, Geofísica e Ciências Atmosféricas-USP (São Paulo, Brasil), 133 p.

França, G.S.L., 2003. Estrutura da crosta no Sudeste e Centro-Oeste do Brasil usando a Função do Receptor. Tese de Doutorado, Instituto de Astronomia, Geofísica e Ciências Atmosféricas-USP (São Paulo, Brasil), 143 p.

Goldstein, P., A. Snoke, 2005. SAC Availability for the IRIS Community, Incorporated Research Institutions For Seismology, Data Management Center, Electronic Newsletter.

Langston, C.A., 1979. Structure under Mount Rainier, Washington, inferred from teleseismic body waves. *J. Geophys. Res.* 85, 4749–4762.

Pavão, C. G., 2010. Estudo de descontinuidades crustais na Província Borborema usando a Função do Receptor. Dissertação de mestrado, Programa de Pós Graduação em Geociências Aplicadas, Instituto de Geociências, Universidade de Brasília, 143 p.

Wessel, P. and Smith, W. H. F., 1995. The Generic Mapping Tools (GMT) version 4.3.1. Technical Reference & Cookbook, SOEST/NOAA, 61 p.

Zhu, L. & Kanamori, H., 2000. Moho depth variation in southern California from teleseismic receiver functions. *Journal of Geophysical Research*, 105, B2, 2696-2980.