



Relationship between porosity, permeability and pore compressibility

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Abstract

Pore compressibility is one of the key properties in reservoir characterization. Such importance is because it influences other properties as porosity, permeability, seismic velocities and rock rigidity. Hence, the knowledge of pore compressibility and the understanding of the mechanism of how it affects the other properties are decisive for several activities in reservoir exploration, management and production.

This work aims to address the influence of pore compressibility on porosity and permeability using experimental data analysis for providing empirical models that relate those properties. The results indicated that power-law models are appropriate for explaining such inter-dependence and bounced that pore compressibility as the key factor to describe the behavior of porosity and permeability of the rocks under external pressure.

Introduction

Pore compressibility is one of the most important petrophysical properties and can influence storage capacity, fluid movement and geomechanical behavior of the reservoir rocks. Tiab and Donaldson (2012) showed that neglecting that parameter can mislead the estimates of volume of oil-in-place. Those authors also addressed the effect of the external pressure in the porosity of the rocks.

Dobrynin (1962) (apud Tiab and Donaldson, 2012) linked pore compressibility to porosity for explaining the variation on that latter, regarding the changes in the confining pressure. Most recently, Silva Jr. et al. (2015) and Oliveira et al. (2016) studied the relationship between pore compressibility and porosity on rocks extracted from outcrops in USA and Brazil. The first work used Mercury Injection Capillary Pressure Porosimetry (MICP) while the second used pore volume measurements through Helium gas expansion technique.

Early works relating pore compressibility and permeability were developed by McLatchie et al. (1958) and Dobrynin (1962). Different from the relationship to porosity, there are few works addressing the relationship pore

compressibility and permeability in the literature up to date (Dobrynin, 1962, Dong et al., 2010).

This work aims to provide new insights of the dependence of porosity and permeability to pore compressibility through the analysis of experimental data on core plugs.

Theory

Geertsma (1957) significantly contributed for the comprehension of the pressure-volume relationship in porous reservoir rocks. He introduced the concept of three types of compressibility: bulk (C_b), matrix (C_m) and pore (C_p). The determination of C_b and C_m uses relatively simple techniques of rock volumetric deformation.

C_p can be obtained through two different approaches. One is to evaluate pore compressibility as a porous pressure function, "simulating" the production process of a reservoir, in which as the depletion occurs, the pore volume is reduced. In those experiments the confining pressure is kept constant. This type of pore compressibility is usually referred as C_{pp} . The second approach is to measure the porous space variation when the sample is submitted to different confining pressures, but keeping pore pressure constant. This type is referred as C_{pc} and is usually associated with the volumetric variation of rock samples when they are brought to surface after coring. Despite the phenomenological differences between those two compressibilities, both are dependent on effective pressure and can be related one to another (Jaeger *et al.*, 2007). Zimmerman *et al.* (1986) re-derived the relationships between different compressibilities and in terms of the confining and pore pressure. The pore compressibility as a function of the confining pressure is defined as:

$$C_{pc} = -\frac{1}{V_p} \left(\frac{\partial V_p}{\partial P} \right)_{P_p} \quad (1)$$

Method

An Ultrapore-300 helium gas expansion Porosimeter was used to measure grain volume and grain density. A digital caliper allowed the measuring of the length and diameter of the cylindrical samples enabling to estimate the bulk volume and then calculate pore volume and porosity. Operational procedures were performed according to API, 1998.

A poroperm instrument (Coreval 700) was used to perform pore volume (PV) measurements and to determine rock permeability at different confining pressures. The pore volume was inferred through gas

(N₂) expansion technique (API, 1998) while permeability was evaluated through an unsteady state technique known as pressure falloff (Jones, 1972).

The same equipment also measured the length of the core plugs at each pressure stage using a linear potentiometric transducer with 0.01 mm accuracy. Those measurements can be used for estimating the vertical strain of the rocks.

Porosity was estimated using Eq. 2 and considering that the grain volume (GV) remained unaltered during those experiments.

$$\phi = \frac{PV}{(PV + GV)} \quad (2)$$

Pore compressibility was estimated through the method described by Unalmsier and Swalwell (1991), which consists in fitting a power-law curve relating the pore volume variation upon external pressure, and estimate C_{pc} using the derivative of that power-law function (Eqs. 3-5).

$$V_p = aP_c^{-b} \quad (3)$$

$$\frac{\partial V_p}{\partial P_c} = -abP_c^{-b-1} \quad (4)$$

$$C_{pc} = \frac{-b}{P_c} \quad (5)$$

Rock Samples

We used a set of 9 rock samples consisting in 6 sandstones and 3 carbonates extracted from outcrops in the USA. Table 1 lists the mineral content of those rocks according to other works in the literature.

Results

Figure 1 shows the variation of the pore volume as function of the confining pressure for the rock samples. A first-order power-law fitting (Eq. 3) is likely to describe very well such behavior presenting coefficients of determination (R²) greater than 0.82. Such values as also the coefficients of the power-law are listed in Table 2.

Pore compressibility was estimated using Eq. 5 and the coefficient b listed in Table 2. A crossplot between those values and the confining pressure is displayed in Figure 2.

Figure 3 shows a comparison between the results of pore compressibility of Berea and Boise sandstones obtained in this work and extracted from other works in literature. That comparison aimed to check how close are our results, which were obtained using gas porosimetry under hydrostatic conditions, to the results obtained with other methodologies. For Berea the results are consistent with Zimmerman (1984) and Oliveira et al. (2015). The first performed C_{pc} measurements on brine-saturated samples by measuring the rock deformation using strain gauges and measuring the expelled brine. The latter used pore

volume measurements obtained through He gas expansion technique, but at uniaxial stress condition. C_{pc} were calculated using the Unalmsier-Swalwell method.

For Boise sandstone, the results are very similar to the reported by Zimmerman (1984).

The relationship between porosity (ϕ) and pore compressibility is exhibited in Figure 4. Such behavior can be expressed through a second-order power-law fitting (Eq. 6).

$$\phi = f(C_{pc})^g + h \quad (6)$$

The coefficients f, g and h are listed in Table 3. Those regressions resulted in R² ranging from 0.976 up to 0.999.

Crossplots between the logarithm of gas permeability and pore compressibility as shown as in Figure 5 allowed evaluating the relationship between those properties. The best fitting resulted in a second-order power-law expression (Eq. 7).

$$\log(k) = m(C_{pc})^n + o \quad (7)$$

The coefficients m, n and o are exhibited in Table 4. Such fitting presented R² ranging between 0.708-0.995.

Discussions

The power-law regressions are likely to fit the behavior of the pore volume due to confining pressure variation with R² ranging from 0.824 to 0.992. Idaho sandstone exhibited the worst fitting. Figueiredo (2015) presented Mercury Injection Porosimetry (MIP) results obtained in similar samples and reported that pore sizes ranging from 10-100 μ m. Possibly, those larger pore sizes are associated to post-diagenetic events. Kahl et al. (2016) showed results of X-ray micro-CT obtained in this kind of rock and reported weathered minerals as feldspar and muscovite. The images presented by those authors showed large non-rounded grains and a variety of pore-shapes. These features may lead to a deviation from the expected power-law trendline in the pore volume-pressure curve.

Porosity is likely to have a great dependence on pore compressibility. A second-order power-law equation perfectly describes such dependence, characterized by the high values of R² (<0.97). Such power-law model also describes the relationship between absolute permeability and pore compressibility, but differently than ϕ -C_{pc} behavior, the R² values were smaller for some of the samples, indicating that other factor may influence permeability rather than solely the pore compressibility on those rocks. Even though, 67% of the samples reported R² greater than 0.9 highlighting the feasibility of that power-law model.

The coefficients of those power-law equations are specific for each sample. Probably they are associated to certain characteristics of the rock, such as mineral content,

unconfined porosity, pore-shape distribution and aspect ratio.

Conclusions

This work showed that such porosity as absolute permeability can be related to pore compressibility through power-law models. The high accuracy of those models indicates that pore compressibility has a major influence on the variations of porosity and permeability when the rocks are submitted to external pressure.

Although the coefficients of those expressions vary for different rocks, perhaps such models may be used for estimating pore compressibility or even absolute permeability from well log data, if calibrated coefficients were available.

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Table 1: Mineral content of the rock samples (in %) and lithologies obtained through other works.

Sample	Calcite	Quartz	Dolomite	Albite	Biotite	Kaolinite	Orthoclase	Plagiocase	Feldspar	Mica	Clinoptilolite	Lithology	Reference
Berea		91				4		1	3			Sandstone	Kokurek (2019)
Boise		49						29	12	5	4	Sandstone	Kokurek (2019)
Indiana Limestone	99.9	0.1										Limestone	Manhães (2018)
Winsconsin	0.7	16.3	83									Dolomite	Figueiredo (2015)
Nugget		72.34			7.91	6.79	12.96					Sandstone	Manhães (2018)
Colton	5.16	64.24	9.22	9.23		12.14						Sandstone	Manhães (2018)
Crab Orchard		93							0.95	6.05		Sandstone	Figueiredo (2015)
Edward Yellow	99.7	0.3										Limestone	Figueiredo (2015)
Idaho		49						29	12	5	4	Sandstone	Kokurek (2019)

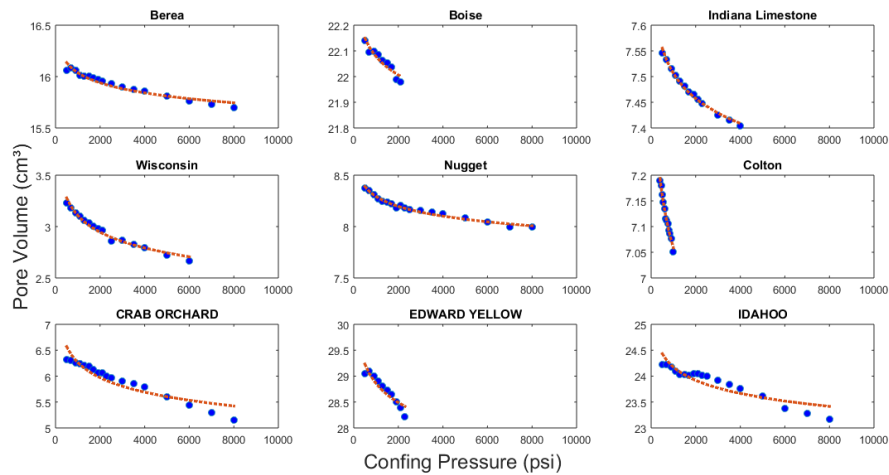


Figure 1: Crossplot between pore volume and confining pressure (blue dots). Red dashed line represents a power-law fitting.

Table 2: Coefficients and R² of the first-order power-law fitting showed on Fig. 1.

Sample	a	b	R ²
Berea	17.048	-0.0089	0.947
Boise	22.803	-0.0047	0.886
Indiana Limestone	8.010	-0.0094	0.992
Winsconsin	5.299	-0.0773	0.977
Nugget	9.315	-0.0168	0.985
Colton	8.145	-0.0207	0.990
Crab Orchard	10.096	-0.0691	0.867
Edward Yellow	32.805	-0.0185	0.850
Idaho	26.877	-0.0153	0.824

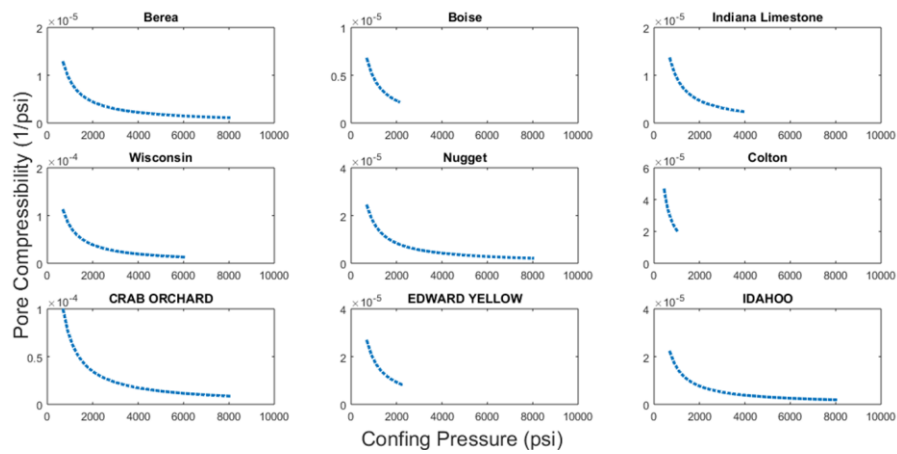


Figure 2: Crossplot between pore compressibility (estimated using Unalmeris-Swalwell technique) and confining pressure.

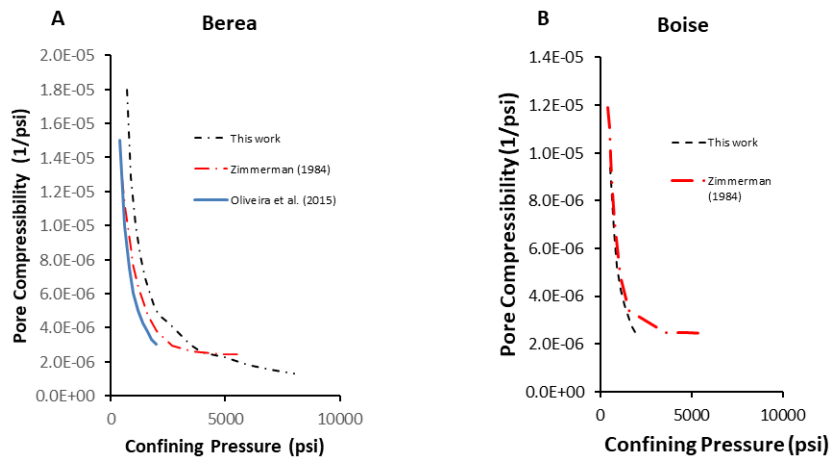


Figure 3: Crossplot between pore compressibility and confining pressure for various works.

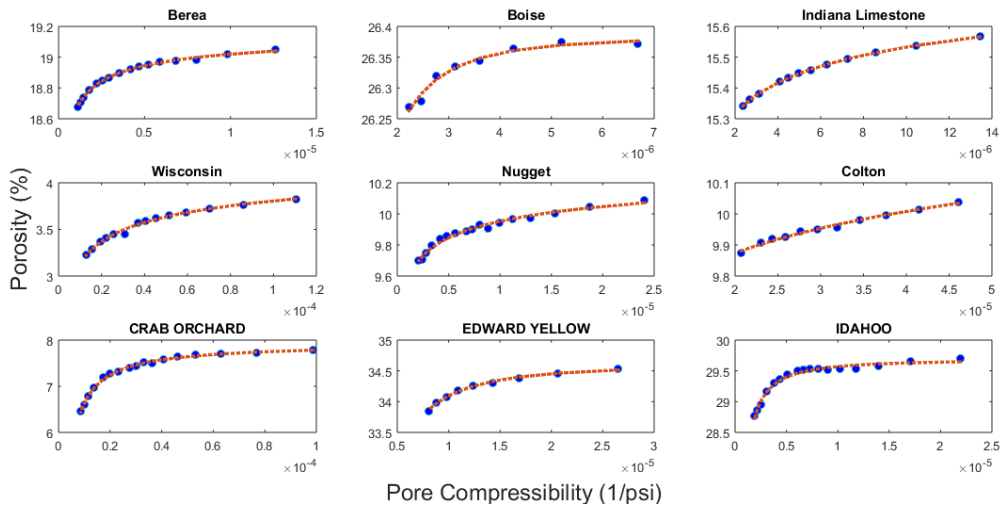


Figure 4: Crossplot between porosity and pore compressibility (blue dots). Red dashed line represents a second-order power-law fitting.

Table 3: Coefficients and R² of the power-law fitting showed on Fig. 4.

Sample	a	b	c	R ²
Berea	-1.53E-03	-0.430	19.236	0.997
Boise	-4.17E-16	-2.559	26.384	0.976
Indiana Limestone	-8.97E-02	-0.174	16.201	0.999
Winsconsin	-1.54E-01	-0.211	4.883	0.992
Nugget	-2.07E-01	-0.140	10.987	0.986
Colton	-7.28E-01	-0.096	11.941	0.989
Crab Orchard	-3.87E-05	-0.906	7.948	0.997
Edward Yellow	-2.46E-09	-1.665	34.619	0.996
Idaho	-8.42E-08	-1.234	29.695	0.983

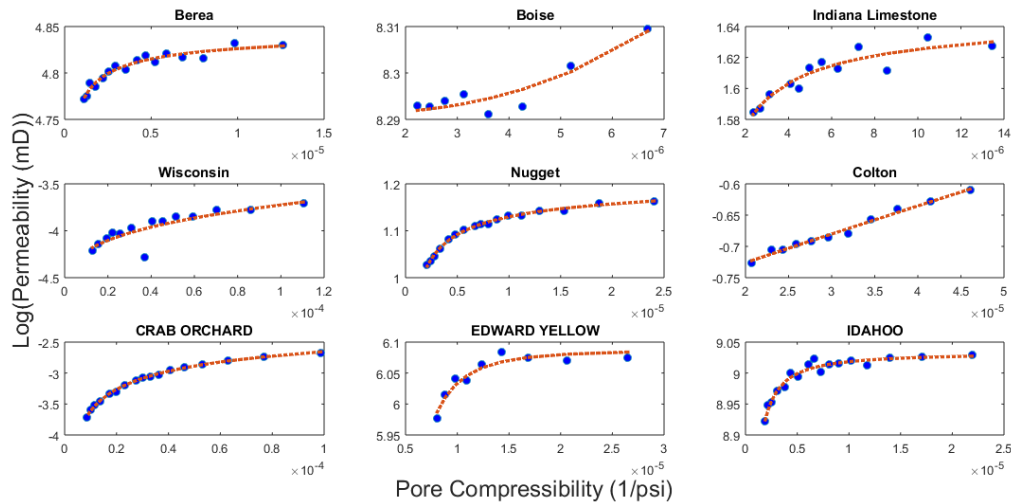


Figure 5: Crossplot between the log of gas permeability and pore compressibility (blue dots). Red dashed line represents a second-order power-law fitting.

Table 4: Coefficients and R² of the power-law fitting showed on Fig. 5.

Sample	m	n	o	R ²
Berea	-7.51E-05	-0.508	4.852	0.948
Boise	9.12E+11	2.648	8.291	0.873
Indiana Limestone	-1.54E-05	-0.649	1.652	0.879
Winsconsin	1.56E+01	0.299	-4.716	0.708
Nugget	-5.65E-04	-0.452	1.233	0.995
Colton	2.65E+03	0.943	-0.824	0.986
Crab Orchard	-1.25E-01	-0.249	-1.418	0.997
Edward Yellow	-1.57E-16	-2.907	6.087	0.917
Idaho	-1.19E-08	-1.217	9.033	0.960