

Annual effective and air absorbed doses based on airborne and ground gamma-ray spectrometry: Serra do Carambeí and Joaquim Murtinho granites, southern Brazil

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Abstract

We investigated the air absorbed dose (AAD) and the annual effective dose (AED) of the Serra do Carambeí (CA) and Joaquim Murtinho (JM) granites, based on airborne and ground gamma-ray spectrometry data. Important relationships are considered between each radioelement (K, $U_{eq}\,$ and $Th_{eq})$ and the effective dose distribution, with effects of topographic and environmental variations. High values of potassium and uranium and very high values of thorium were confirmed. The mean average of AAD is 100% (airborne) and 170% (ground) in CA and 35% (airborne) 218% (ground) in JM upper the global mean average. Both granites showed values of AED superior to the recommended by international health organizations. The CA granite revealed approximately twice the safety value in both data types. For the JM granite the values are almost five times (ground data) and about one and a half times greater (airborne data) than those suggested for outdoor exposure.

Introduction

Several epidemiological studies have been designed to assess hazards from exposure to natural radiation (Grastly & LaMarre 2004, Ye & Gan 2013, Oatway et al. 2016). A large effort is being made by nations to understand and minimize the consequences of this that, although it is a natural process, it can entail damages to human's health when background radioactivity reaches high levels (US EPA 1993; EC 1999; UNSCEAR 2000).

Radioactivity on Earth is a process that occur constantly and has terrestrial and extraterrestrial origins, being part originated in subsurface by decay of the parent nuclides ²³⁸U, ²³²Th and the nuclide ⁴⁰K. Upon decay, these radionuclides produce an external radiation field to which all individual are exposed being them the principal dose radiation source on earth (UNSCEAR 2000). The measurement of radiation from rocks and soils are made by detection of gamma-rays, that identifies each radioelement through their respective energy spectrum. In this way, we can estimate dose rates, radon concentration and other parameters from readings of K (%) and equivalents of U (ppm) and Th (ppm) in rocks and soils up 30-45 cm deep (Dickson & Scott 1997; Minty 1997; Wilford et al. 1997; IAEA 2003).

The earth's upper crust has about 2-2.5 % of K, 2-3 ppm of U and 8-12 ppm of Th (Ulbrich et al. 2009). The average of

these two radioelements are 15 ppm for Th and 4 ppm for U in igneous rocks with granitic composition. The potassium content is relatively low in granites (~2.4%) and higher in felsic volcanic and pegmatites (~3.7%). In granites, the concentration of Th and U is higher when compared with basaltic or ultramafic rocks (< 1 ppm and < 8 ppm for U and Th, respectively) (Ulbrich et al. 2009).

The radioactivity on rocks depends on their original composition when formed during the molten. However, the radiometric signature can vary a lot even over the same formation due changes in content of radioelements by small lithological differences, weathering, regolith effects and mineralizing events (Dickson & Scott 1977). On air, the presence of soil-derived radionuclide will depend of local rocks and soil, wind, and moisture conditions (UNSCEAR 2000).

The global mean value of the dose in air according to the distribution of the population ranges 50–59 nGy h⁻¹ (UNSCEAR 2010, Engola et al. 2018). The detriment per unit of effective dose for public members is 7.3 10^{-5} mSv, being 1000 µSv the limit accepted in a year. Higher values can be considered a hazard by the United States Environmental Protection Agency (US EPA 1993) and by the Organization for Economic Co-operation and Development (EC 1999).

The scope of this work is to estimate the AAD and the AED of the CA and JM granites, based on the airborne and ground measurements of K (%), U_{eq} (ppm) and Th_{eq} (ppm), in order to assess the risks associated to the human health. The results are compared with indices suggested by international medical organizations and radiological institutions.

Study Area and Geology

Related to Upper Proterozoic and after the collisional events of Brasiliano Cycle, Serra do Carambeí (CA) and Joaquim Murtinho (JM) are two granitic bodies that they're part of the Cunhaporanga Granitic Complex (Guimarães 2000), which are localized in State of Paraná, southern Brazil (Fig. 1). Both are equigranular and fluorite is one of secondary minerals.

The CA granite (e.g. Pinto-Coelho 1986; Guimarães 2000) has an elongated form in NE-SW direction, with 14 km length and 7.3 km width (Fig. 1), and it is a biotite alkali feldspar granite with quartz and hematite. Fluorite and zirconium are secondary minerals.

The Joaquim Murtinho granite (CCGP 1970; Guimarães 1995; Guimarães & Ulbrich 1996) is an igneous intrusion with irregular shape and area < 10 km² (Fig. 1). It forms a relative local depression, delimitated by sandstones of Furnas Formation in east, north and west, monzogranite in south and alluvial deposits at south-west. It is composed mainly of alkali feldspars, biotite and quartz. The fractures

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are filled with hematite, fluorite and/or sulfides (Mineropar 2006; Ulbrich et al. 2009).



Figure 1 – Location and geological units of the Serra do Carambeí (CA) and Joaquim Murtinho (JM) granites (modified from Mineropar 2006).

Material and Methods

Data

We used airborne and ground gamma-ray spectrometry data (K, %; Th_{eq} , ppm; U_{eq} , ppm) to estimate the Air Absorbed Dose (AAD) and the Annual Effective Dose radiation (AED) of the Serra do Carambeí (CA) e Joaquim Murtinho (JM) granites. The database comprises 850 readings in the area CA granite: 766 airborne measurement and 84 terrestrial stations. Inside the limits attributed to the JM granite were read 348 points: 310 from airborne survey and 38 measurements on ground (Tables 1 and 2).

The airborne data are from Paraná-Santa Catarina Project (CPRM 2011). The flight lines in NS direction was spacing of 500 m, with a mean terrain clearance of 100 m and the time interval between the measures was 1 second.

The authors Ferreira et al. (2001) and Fruchting et al. (2003) firstly presented the ground data of the CA and JM granites, respectively. Detailed discussion about the

terrestrial data of JM and CA can be seen in Ferreira et al. (2009) and Godoy et al. (2011), respectively. The ground data was acquired by a GS-512 gamma-ray spectrometer, manufactured by Geofysika from Czech Republic. Tables 1 and 2 and Figure 2 shows, respectively, the basic statistics and the graphs of the CA and JM data.

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Radiation doses estimates

Radiation dose estimates are possible based on the K (%), U_{eq} (ppm) and Th_{eq} (ppm) measurements. The International Atomic Energy Agency (IAEA 1991) and the International Commission on Radioactive Units of Measurements (ICRU 1994) recommended an air absorbed dose (AAD) estimation based on airborne gamma-ray spectrometric surveys (Ye & Gan 2013, Beamish 2016), according to equation (1):

AAD (nGy
$$h^{-1}$$
) = 13.078 K + 5.675 U_{eq} + 2.496 Th_{eq} (1)

The constants that multiply K (%), U_{eq} (ppm) and Th_{eq} (ppm) in the equation (1) were suggested by IAEA (2010) as default parameters. 1 Gy is equal to 100 rad or 1 J/kg.

Dose rate indicates the quantity of radiation absorbed per unit of time. A similar measurement is radiation dose, which refers to the rate with which a person has been exposed to a radiation source.

The sum of multiples of equivalent doses in human organs is measured by the effective dose, where each organ has a specific weighting factor. The adopted form to evaluate the contribution of natural radionuclides only from soil, rocks and manufactured materials during a year, excluding extraterrestrial radiation, is the annual effective dose radiation (AED). AED quantifies the biological and medical effects associated with the type of ionizing radiation during a year (Ray & Stick 2015, Beamish 2016). Godoy et al. (2011) proposed the equation (2) based on Grasty et al. (1984).

AED = $320 + 52.6 (1.505 \text{ K} + 0.625 \text{ U}_{eq} + 0.310 \text{ Th}_{eq}) (2)$

The SI unit adopted by IAEA (2003) is the micro Sieverts per year (μ Sv y⁻¹), the most used for studies that connect radiation effects in human body.

The AED expression considers the contribution of K, U_{eq} and Th_{eq} to the exposed rate at 1 meter above the ground and the cosmic radiation from a nearby site (Pascholati & Amaral 1992). The parameters of the equation 2 take in to account the absorbed dose in air and the effective dose utilizing the indoor–outdoor occupancy factor (UNSCEAR 2000), where the numerical values vary with the age of the population and the local climate (Beamish 2006).

Results and Discussion

The airborne and ground data were interpolated by Minimum Curvature method (Briggs 1974), and the contour maps superimposed on the Digital Elevation Model (DEM) are presented in Figures 3 and 4. The shapefiles of both granitic bodies are from Mineropar (2006).

Analysis of Serra do Carambeí (CA) granite

Potassium and uranium presented discrete values in the airborne data when compared with the respective crustal averages (Tab. 1). Highest contents of K occur on valleys and at bottom of slopes in the alkali feldspar granite (Fig. 3). This behavior is due to the effects of erosion, depletion, transport and deposition of K that cumulates in lowest zones and nearly downhills. In high altitude regions and low declivity, the concentration of U is not expressive. This can be explained due the effect of percolation of the water in plane areas that acts in favor of lixiviation of rocks and transport of K and U⁺⁶. The high mobility of K explains the lower concentration on the top of CA granite. Concentration of $U_{eq}\,$ and Th_{eq} are very similar and occur most part in the north and along the east, circulating and delimiting the body. Uranium and thorium are most located on the edges of the granite, but uranium particularly showed bigger area on south. In the greatest depressions, where ticker layers of sediments, organic matter and soil is present, the uranium concentration shows low (Fig. 3).

Table 1. Statistic radioactivity measurements of CA

	Airborne			Ground		
	К	U _{eq}	Th _{eq}	К	U _{eq}	Th _{eq}
samples	766	766	766	84	84	84
mean	0.53	4.47	28.44	1.66	7.25	39.08
median	0.46	4.18	27.32	0.30	6.10	38.55
std	0.36	1.76	10.39	1.89	3.92	15.23
min	0.00	0.57	11.84	0.10	1.20	6.40
25%	0.29	3.33	22.39	0.10	4.60	29.37
50%	0.46	4.18	27.32	0.30	6.10	38.55
75%	0.69	5.28	31.22	3.40	8.57	48.20
max	2.18	12.95	90.71	6.10	19.20	82.20

Table 2. Statistic radioactivity measurements of JM

	Airborne			Ground		
	К	U _{eq}	Th _{eq}	К	$\mathbf{U}_{\mathbf{eq}}$	Th _{eq}
samples	310	310	310	38	38	38
mean	1.86	4.36	17.13	3.8	9.22	40.68
median	1.31	3.46	10.88	4.65	6.65	29.35
std	0.60	1.24	8.37	2.36	8.10	37.98
min	1.10	2.23	7.46	0.10	1.40	8.90
25%	1.37	3.47	10.34	1.02	5.25	22.17
50%	1.72	4.27	14.32	4.65	6.55	29.35
75%	2.16	5.02	22.42	5.77	8.60	37.50
max	3.50	8.10	49.76	7.50	38.00	161.50

It is must be noted that CA presented almost ten times more Th_{eq} than the crust average. High concentration of Th_{eq} and U_{eq} occur with more frequency in areas of high elevation, but we verify a distribution of highest values across east and north. Therefore, as expected, the topography was not a decisive factor to the thorium behavior. On the other hand, we verify "hot" areas where the topographic gradient is accentuated, probably because in these parts, the rock is more expose and sediments or soils cannot mask the radiometric original response.

Ground and airborne data presented different spatial distribution patterns but compatible values (Fig. 3). In the ground data, the K mean is about four time higher than the airborne one (0.53% – 1.66%), while the maximum concentration of U_{eq} increases by 6.25 ppm (12.95 - 19.20 ppm). The maximum ground concentration of T_{heq} is 8.5 ppm lower than the airborne correspondent data (Fig. 2).



Figure 2 – Graphs of K, U_{eq} and Th_{eq} concentration of (a) Serra do Carambeí and (b) Joaquim Murtinho granites. Black thin bars represent the standard deviation.

The mean estimates doses of AAD and AED are in the Table 3 and 4. As we can see, the ground data is more disperse demonstrating the wide range of values acquired by this type of survey.

Granite			AAD (nGy h	⁻¹)
	Туре	Mean	std	Sample
Carambeí	Airborne	103.30	33.07	766
	Ground	152.80	42.75	84
Joaquim	Airborne	73.29	25.85	310
Murtinho	Ground	171.07	115.14	38

Table 4. Estimates of annual effective dose (AED)

Granite			AED (µSv y	-1)
	Туре	Mean	std	Sample
Carambeí	Airborne	972.65	225.21	766
	Ground	285.31	383.54	84
Joaquim Murtinho	Airborne	889.69	194.30	310
	Ground	1589.77	930.65	38



Figure 3 – CA gamma-ray spectrometric maps of K (a), U_{eq} (c), Th_{eq} (e), AAD (g) and AED (i) from airborne data and K (b), U_{eq} (d), Th_{eq} (f), AAD (h) and AED (j) from ground

data. The isolines of 1000 $\mu Sv~y^{-1}$ are insert in the AED maps.

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The maps of the AED of CA (Fig. 3 i,j) and AAD (Fig. 3 g,h), shows that the annual effective doses are affected mainly by U_{eq} and Th_{eq} content, being thorium the major influence. The highest values of AED are seen in the east edge and north of the granitic body and in more elevated areas with values up to 2229 µSv y⁻¹ (airborne) and 2146 µSv y⁻¹ (ground). The highest values of AED are seen in the edge and north of the granitic body and in more elevated areas with values up to 2229 µSv y⁻¹ (airborne) and 2146 µSv y⁻¹ (Ground). The maps of AAD and AED have the same characteristics in terms of distribution. The data collected on ground present more dispersive pattern than that of collected on air. AAD showed mean values up the 100 nGy h⁻¹ and a maximum of 300 nGy h⁻¹ (Tab. 3, Fig. 4 g,h), much higher than the 51 nGy⁻¹ recommended.

Analysis of Joaquim Murtinho (JM) granite

All three radioelements of the ground data processing showed anomalous high zones that overcome the crust average, but in the airborne data only Th revel elevated counting, reaching almost 200 ppm (Fig. 4). None of them shows strong correlation between the geophysical anomalies and topography, this means that the concentration of K, U and Th seems more related to lithologic itself and the different facies of JM granite. Even so, it is possible to observe that the anomaly of U and Th set in the highest portion of the map whereas the K present a more fluid behavior, allocated and distributed on slopes with high gradient variation but near to the same U and Th anomaly. In the center (Fig. 4 a,b,c,d,e,f) strong anomalies of U_{eq} and Th_{eq} associated with depletion of K, suggest an accentuated process of weathering. Highest values of Potassium concentrate in south-west and a small part in the south and center-east of JM. In general, the body present medium-strong K, associated with the feldspathic quartz composition of the igneous rock. Uranium shows very low values in the east, where the topography is compost by valleys and abrupt slopes. The principal Ueq anomaly (Fig. 4 c,d) covers the center-southeast of the JM granite. In same location, a strong anomaly of Theg shows content around 35-50 ppm in the airborne data and about 100-200 ppm in the ground data (Fig. 4 e,f).

To all radioelements, from minimum to maximum and including the mean and median, the terrestrial data presented much higher values in comparison than airborne data (Tab.2, Figure 3).

The mean of AAD of JM in the terrestrial data was 100 nGy h⁻¹ higher of that found in the airborne database, being the maximum 4 times higher. There was no measure below Again, Th influences more the AED distribution than potassium and uranium on both data type, being Th responsible for the highest percentage. AED attained almost 1500 μ Sv y⁻¹ (airborne) and 4931 μ Sv y⁻¹ (ground) in the light pink spots. This difference between AED calculation is connected to the much higher reading of Th_{eq} on ground when compared with the maximus content of the same element in the airborne data (Fig. 4 e,f), with maximum reading of Th_{eq} overtaking 150 ppm where the maximum is 50 ppm in the aero survey data (Tab. 2).



Figure 4 – JM gamma-ray spectrometric maps of K (a), U_{eq} (c), Th_{eq} (e), AAD (g) and AED (i) from airborne data and K (b), U_{eq} (d), Th_{eq} (f), AAD (h) and AED (j) from ground data. The isolines of 1000 μ Sv y⁻¹ are insert in the AED maps.

Conclusions

Air absorbed dose (AAD) and annual effective dose (AED) radiation reached elevated values being the maximum reading of AED to both granites very high and worrying. The highest value of AED in the CA granite is more than twice the value recommended by IAEA, being 2229 μ Sv y⁻¹ (airborne) and 2146 μ Sv y⁻¹ (ground), with averages of 972 μ Sv y⁻¹ (airborne) and 285 μ Sv y⁻¹ (ground). In the JM the AED obtained almost 1500 μ Sv y⁻¹ (airborne) and around

5000 μ Sv y⁻¹ (ground), with averages of 889 μ Sv y⁻¹ (airborne) and 1589 μ Sv y⁻¹ (ground) while the AAD average in CA is 100% (airborne) and 170% (ground) above the recommended safety value and 35% (airborne) and 218% (ground) above in JM.

The Th seems to be the radioelement less affected by the weathering. The high measurements of Th_{eq} influenced more the results of AED and AAD than K and U_{eq} .

The spatial distribution of K, Th_{eq} and U_{eq} have been studied through airborne and terrestrial surveys. The ground data presented higher values than the aerial ones, probably due the different sampling types (point sampling versus field of view sampling, respectively), wind and because the radiation detected is stronger near the ground, suggesting that important attenuating processes could occurred in air between soil and detector. In addition, portion of gamma-ray flux must have been affected by variation on soil thickness, moisture, organic matter and lixiviation processes, so we cannot link the variation of doses only to the composition of the granitic rocks, since it is known the importance of soil characteristics to radiation ionization.

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