

Geophysical modeling of Serra Negra and Salitre I, II and III alkaline-carbonatite complexes based on their gravimetric and magnetic signatures

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Abstract

The alkaline complexes of Serra Negra and Salitre I, II and III are part of the Alto Paranaíba Igneous Province (APIP), which has been subject of many mining projects due to its important content in mineral resources as phosphates, niobium, titanium, among others. The gravimetric and magnetic anomalies caused by the complexes displays mutual interference becoming difficult to suit individual interpretations. This paper shows the results of 3D modelling of the anomalies and compares the results with the geological known information.

Introduction

The Alto Paranaíba Igneous Province (APIP) (Fig. 1) is an assembly of mafic to ultramafic and potassic to ultrapotassic magma intrusions that are aged in the Late Cretaceous (Campos and Dardenne, 1997). The rocks of the complexes include kamafugites, kimberlites, peridotites, lamproites, alkaline-carbonatite with phlogopite-picrites, dunites, bebedourites, sienites, carbonitites and phoscorites disposed in pipes, plugs, diatremes, lava flows and plutonic intrusions (Gibson et al., 1995; Marangoni and Mantovani, 2013).

The alkaline-carbonatites of Araxá, Tapira and Catalão I and II that belongs to the APIP's province are already been exploited for then mineral resources of phosphates and Niobium and have known deposits of Titanium and REE (Brod et al, 2004). Many studies of complexes of Serra Negra and Salitres I, II and III have been produced in the last few years to understand their origin, evolution and geometry as well as their mineral resources (Grasso, 2010; Uliana, 2010).

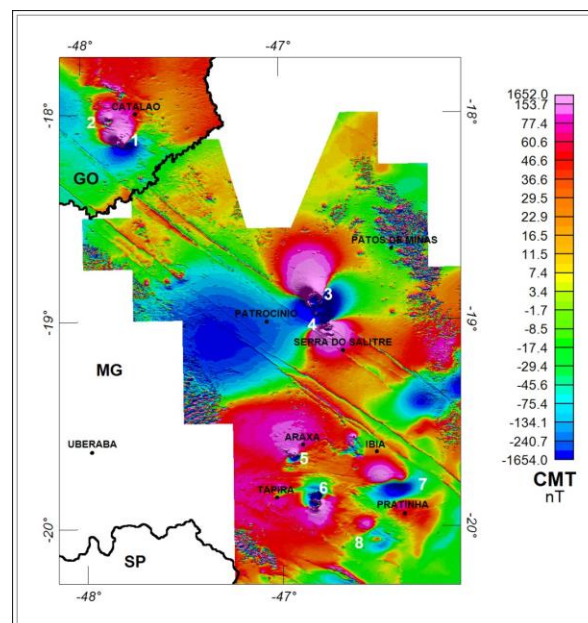


Fig. 1. Total magnetic field anomaly map (nT) for the APIP. Numbers in white indicate alkaline complex locations: 1-Catalão, 2-Catalão 2, 3-Serra Negra, 4-Salitre I, II and III, 5-Araxá (Barreiro), 6-Tapira, 7-Pratinha 1, 8-Pratinha 2. Regional importante cities and state borders are marked in black (extracted from Marangoni and Mantovani, 2013).

The gravimetric and magnetic signatures of the complexes are very distinctive compared to the host rocks and when studied jointly can reduce the amount of ambiguity which is inherent to the potential field methods. The goal of the present study is to compare the results from different approaches by inversion and direct modelling using the magnetic and gravimetric data available over the target and evaluate them with the geological knowledge of the complexes.

Geological Context

The Serra Negra and Salitres complexes are located close to Patrocínio city at Minas Gerais state. Both of them are intrusions in the Canastra group formed of quartzites and metasedimentary rocks.

Serra Negra has a circular shape of approximately 10 Km of diameter. Mariano & Marchetto (1991) describes the complex as formed by Calcium Carbonatite in his center with apatite, dolomite and minor parts of phlogopite,

magnetite, humite, rutile, pyrochlor, pyrrhotite and pyrite. Amaral et al. (1967) showed ages between 83.7 and 83.4 Ma for the complex. The geological map of Serra Negra is shown in Fig. 2.

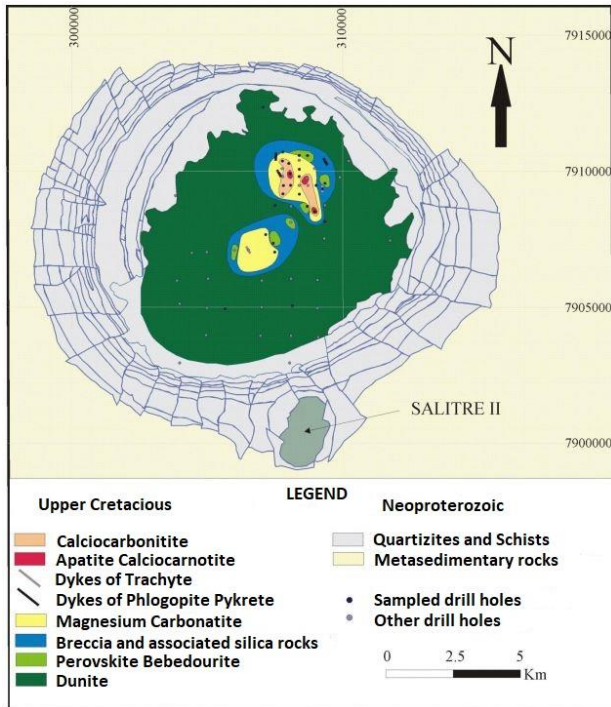


Fig. 2 – Geological Map of Serra Negra alkaline complex modified from Grasso (2010).

Barbosa et al. (2012) describes the Salitre I alkaline complex as a plutonic intrusion with the lithotype predominance of Bebedourites, Phoscorites and carbonites. According to Sonoki & Garda (1988) the K/Ar dating at phlogopites in Bebedourites gives the age of 86.3 ± 5.7 Ma for the complex. The geological Map of Salitre I is shown in Fig. 3.

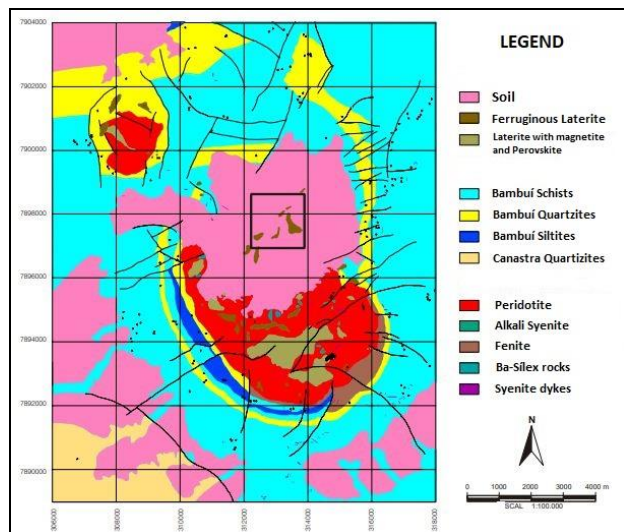


Fig. 3. The Geological Map of Salitre I and II modified from Uliana (2010).

Salitre III is described as small occurrence south to Salitre II with approximately 1,7 Km of diameter. According to Dnmp and Fosfertil presentation for the mining prospects it has Bebebourites and Laterites of Bebedourites as it's major components. The geological map of Salitre III is presented in Figure 4.

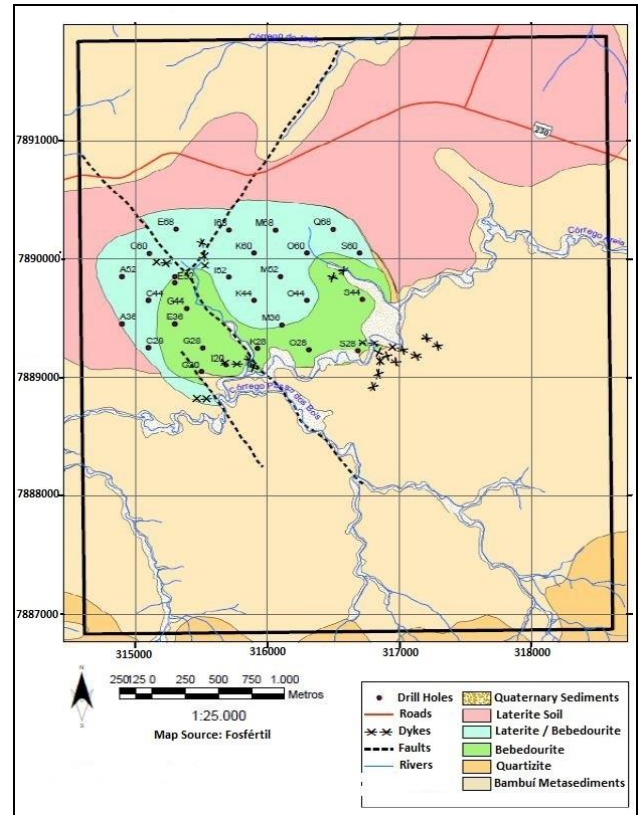


Fig. 4 – Geological Map of Salitre 3 alkaline complex modified from Fonseca (2009).

Method

Magnetometry

The magnetic data in this study was the geophysical airborne survey Área 7 from CODEMIG/CPRM (Brazilian geological agencies) flown between 2005 and 2006. The survey lines have north-south direction and are 400m apart. The mean survey terrain clearance was 100m. The magnetic data was collected at 10Hz sample rate. The Total Magnetic Field were gridded using the minimum curve method with the cell size of 1/4 of the line space separation (100m). The final data were levelled and microlevelled for noise removal and the International Geomagnetic Reference Field (IGRF) extracted. The main database was cut to the study area over the anomalies (Fig. 5).

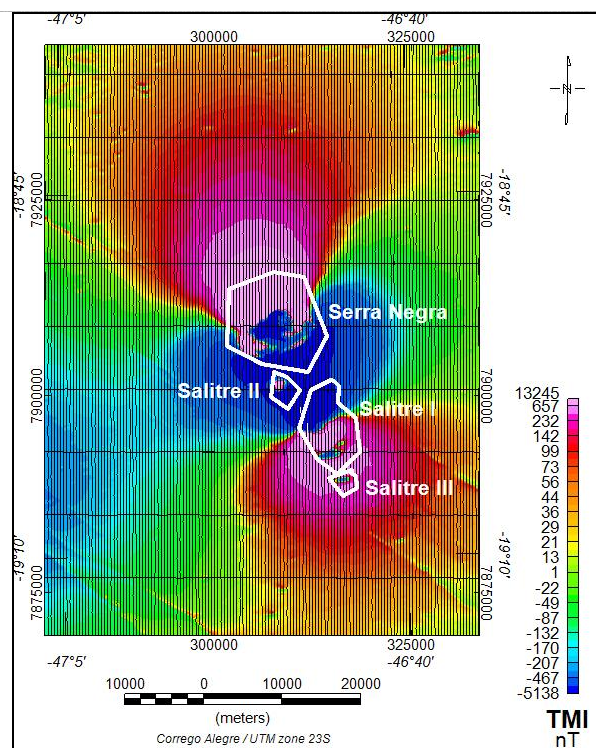


Fig. 5. The residual magnetic field from Serra Negra and Salitre I, II and III. The flight lines are in black.

Forward Modelling

The magnetic forward modelling for the anomalies were obtained using the Modelvision software which assembles a algorithm based on Barnett (1976). The analytical method developed proposes the representation of any arbitrarily polyhedron bodies as composed by a number of finite triangular facets and computes the magnetic field for each facet. The field expression for the whole polyhedron body will be the sum of the each triangle contribution. The software also have the flexibility to use any magnetic information for the bodies as the magnetic remanence or anisotropy.

Inversion Modelling

The inversion modelling for the anomalies were achieved using the Magnetization Vector Inversion (MVI) technique wich is part of Oasis Montaj software. The MVI was presented by Ellis et al (2012) and incorporate remanent and induced magnetization in the inversion algorithm. The autors uses equations between the magnetization components and magnetic anomalies and then optimized the objective function to obtain three components of magnetization vector with the same regularization parameters used in scalar inversions.

Gravimetry

The ground gravimetric data used in this study were acquired by IAG/USP for various projects between 2001 and 2012 totaling 348 stations over all the alkaline complexes. it was used the Lacoste-Romberg model G gravimeter for the acquisition of gravity data. The altimetry was obtained using gps with differential correction allowing precision at about 0.1m. The Bouguer anomaly were calculated using $2,67 \text{ g/cm}^3$ as density background. The grid was generated using minimum curvature with 300m cell size. The residual Bouguer was calculated using The Bouguer anomaly map is presented at Fig. 6.

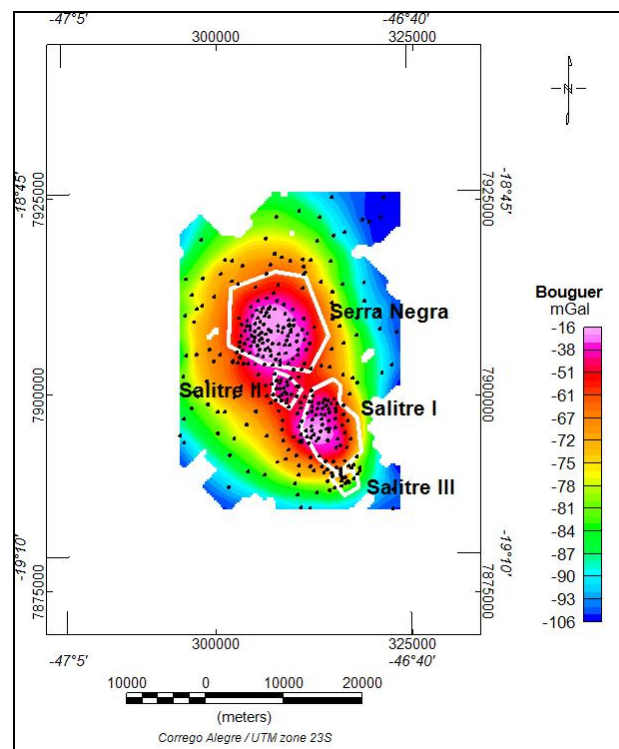


Fig. 6. The Bouguer anomaly map from Serra Negra and Salitres I, II and III. The gravimetric stations are the black dots.

Inversion Modelling

We have used the UBC-GIF GRAV3D software to make the inversion of the gravimetric data over the anomalies. The algorithm is based on Li and Oldenburg (1998), wich bases the inversion as a optimization problem of a objective function, in this case, density contrast, that is minimized for determined constraints reproducing the observed data over an error tolerance.

Results

MAG/GRAV Forward Modelling

For preliminary to model, the Bouguer anomaly grid was sampled to the aeromagnetic survey, so we could use the same database for both methods. Using the amplitude of the analytic signal of the TMI and the geological maps, we

contoured the borders of Serra Negra and Salitre's bodies and try to fit the major shape of the combined anomalies. We have supposed all bodies as plugin prisms and varying its physical properties to adjust the models. The starting remanent magnetization for the reversed anomalies was extracted form Louro (2013). The proposed densities were extracted from Telford et al. (1992).

The final result have considered the time consuming of modelling of all possible magnetic anomalies observed in the profiles, the adjustment between modeled and observed data over the several lines that crossed the bodies and the residual grid from the model and observed data.

The models are presented in the Figures 7,8 and 9.

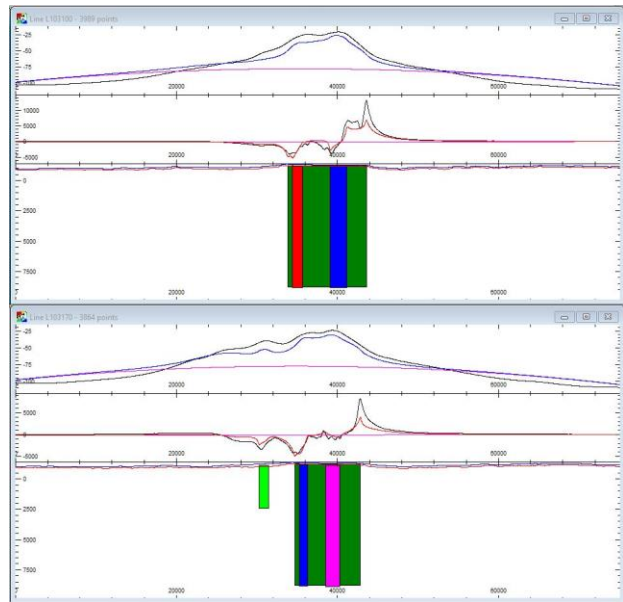


Fig. 7. Flight lines 103100 and 103170 showing the gravity (top) and magnetic (middle) profiles over the suggested anomaly bodies from Serra Negra and Salitre II. The inner bodies were based on the geological maps and observed anomalies within the complexes.

The final proposed density and magnetic values for the bodies differs from the initials as we insert more inner bodies within the complexes in order to better adjust the profiles.

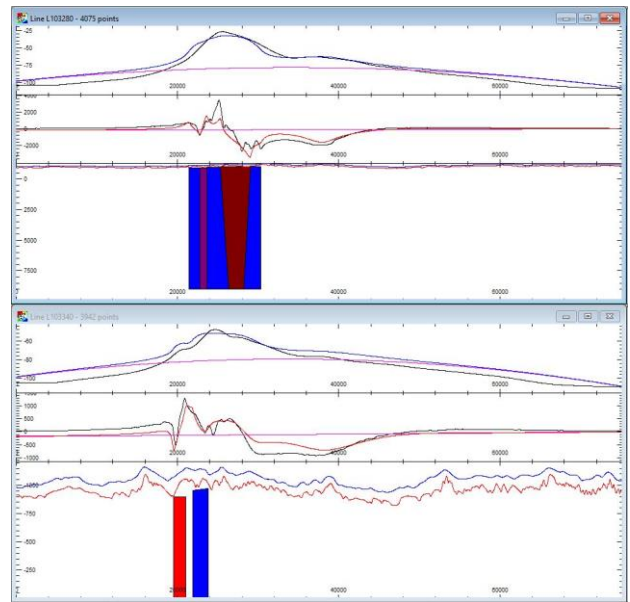


Fig. 8. Flight lines 103280 and 103340 showing the gravity (top) and magnetic (middle) profiles over the proposed anomaly bodies from Salitre I and Salitre III. The inner bodies were based on the geological maps and observed anomalies within the complexes.

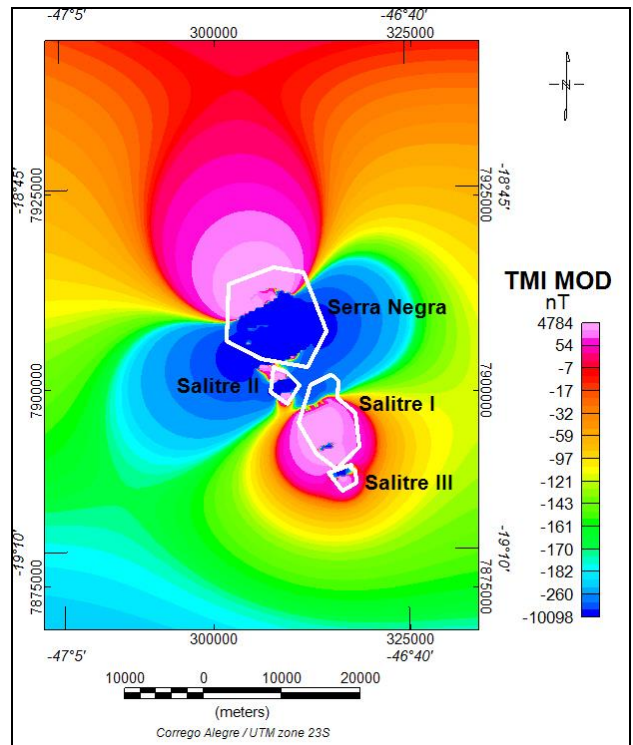


Fig. 9. The total magnetic intensity map from the forward modelled bodies.

The MVI inversion

We used the same database for the MVI inversion. We have chosen 400m for the horizontal cell size and 150m for the vertical in the voxels mesh. For the initial inversions we have not put or changed any constraints or parameters. In the subsequent inversions, we changed the radius of influence, the error attributed to each data and IRI focus as discussed by Pereira et al (2013). The final results were considered very satisfied as it shows clearly the reversed vectors of the magnetization from the anomalies and the shape of the bodies are according to the geological perspective for the alkaline complexes. The amplitude of the apparent susceptibility and the shape recovered from the inversion were also reasonable with a 0.5 SI average for all complexes together.

The model results from the MVI are shown in the figures 10 and 11.

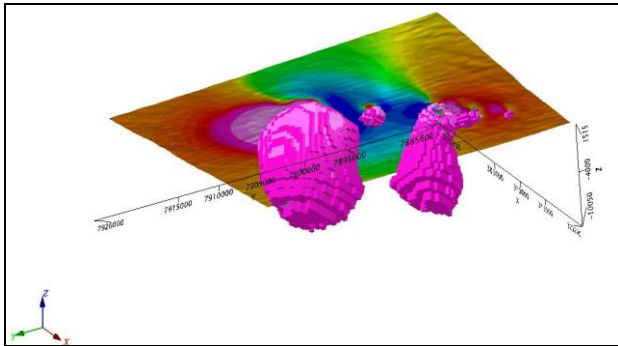


Fig. 10. The 0.07 SI apparent susceptibility clipped surface showing the Serra Negra and Salitre I, II and III bodies.

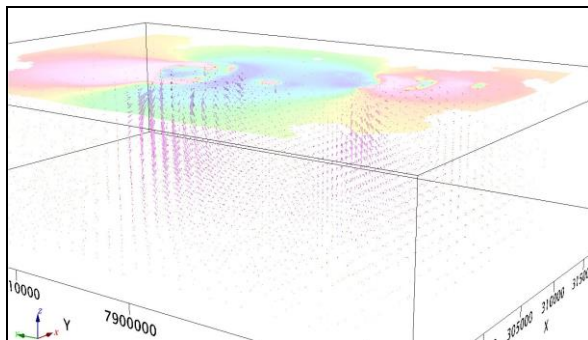


Fig. 11. The magnetization intensity vector recovered from the MVI inversion showing the different magnetization directions for the complexes, including within the same as shown in Serra Negra complex.

Gravity Inversion

In inversion in the GRAV3D software, we have used 348 gravity stations acquired irregularly spaced all over the complexes. The residual Bouguer anomaly was calculated considering the regional field as just a constant plane of 100 mGal. We started the inversions using the GCV trade off to the misfit parameter and used the final results for the subsequent inversions. The mesh cells had 500m horizontally and 100m in vertical. After each

inversion, we have evaluated the results and changed some parameters described by UBC-GIF workflow internet page trying to refine the model. The final result was considered good as it is in accordance with the previous methods applied and the geological available information. The gravimetric models obtained are presented in the figures 12 and 13. The final results are shown together in the same mesh to allow the simultaneous visualization in the figure 14.

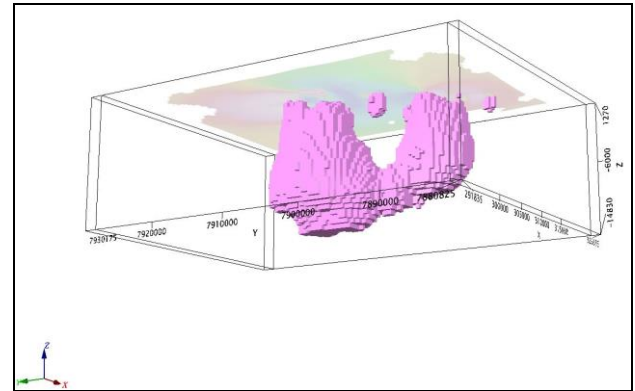


Fig. 12. The 0.3 g/cm^3 recovered density contrast clipped surface showing the Serra Negra and Salitre I, II and III bodies.

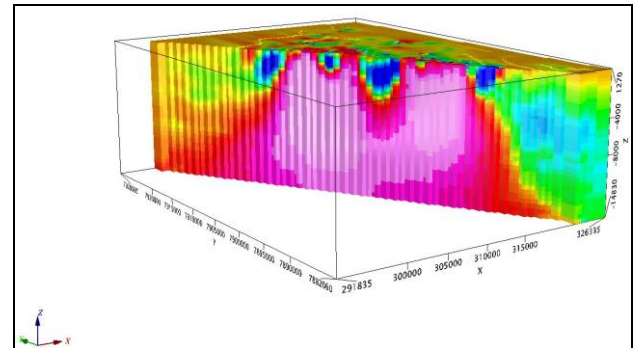


Fig. 13. The amplitude of density contrast sliced at N340W to show the inner part of the models recovered.

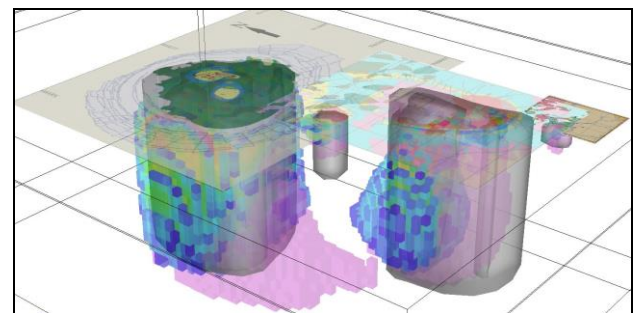


Fig. 14. The three different models in the same mesh. The magenta representing the gravity inversion, the blue to yellow representing the magnetic inversion and in grey the forward magnetic modelling and geological maps.

Conclusions

In this study we have presented three different techniques for modelling the magnetic and gravimetric signature from Serra Negra and Salitres I, II and III. All the procedures showed good results when compared with previous studies of the complexes. The forward modelling demand more time, however it allows the user to get more skills over magnetic signatures and their parameters such as remanent magnetization, geometry of the bodies, etc. The inversions software's used in this study required less processing and time consuming, nevertheless resulted in very good models that are all according to the geological perspectives for the alkaline complexes. The MVI has proved to give an excellent tool to magnetic anomalies that have remanent magnetizations.

Acknowledgments

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