

Application of high-resolution airborne EM and magnetic data in a mining environment for geotechnical information, examples from the Quadrilatero Ferrifero area.

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Abstract

We present an overview of an ambitious geophysical program undertaken by Vale to characterize both tailings storage facilities (TSF) properties around existing mine sites as well as an exploration program for iron ore. Historic magnetics data, geological surveying, borehole information and newly-collected RESOLVE® helicopter frequency domain electromagnetics (EM) and magnetic gradient data is being modelled with 3D joint inversion. We outline the survey goals, data collected, and inversion modeling methodology and show some of the emerging results.

Introduction

There are a large number of mines in Minas Gerais where technical approaches and surviving records are not at the same standard as at modern mines. In some cases, there are no records regarding dam construction. Geophysics presents a non-invasive, cost-effective method of characterizing these dams. However, in places, ground geophysics is not permitted because of HSE concerns. Also, given the number of mine sites, ground geophysics would be an onerous endeavor.

Airborne geophysics is thus a compelling solution. There are two main reasons for airborne acquisition: (1) HSE for conducting ground surveys on top of the existing earth dams and tailing ponds and (2) speed of acquisition. To properly characterize the site, data must be collected in multiple lines over the dam, pond and the immediate surrounding area. We need more than just highly-localized information from ground-resistivity; we need

project-scale data about geologic conditions. There are dozens of areas and thousands of line-kms of acquisition to be accumulated. Speed is important because decisions have to be made on maintenance and to ensure proper and safe operation.

To characterize the dam, very near surface information is required. Frequency domain EM presents the best available near-surface sensitivity of airborne geophysical techniques. Xcalibur's RESOLVE® system has the highest available operation frequency at 140 kHz and has been selected for this project. In order to enhance sensitivity to magnetic material in the ground, RESOLVE was flown with a transverse horizontal magnetic gradiometer.

To ensure the correct interpretation of the complete earth response in these topographically complex settings, the data is being modelled with a 3D inversion program. A cooperative joint inversion is being conducted between the magnetics and the EM, and where available, surface, geotechnical and borehole information is being incorporated..

19,000 line-km will be flown over 23 blocks in the Quadrilátero Ferrífero region, Minas Gerais and Carajas, Southern Para. Line spacing ranges from 25m to characterize the dam sites with high resolution to 100m over the exploration blocks. Extensive culture (buildings and power lines) caused the survey to be flown 30 m higher than typical surveys.



Figure 1 The surveys around Belo Horizonte.

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The Resolve FDEM System

Figure 2 shows a picture and schematic diagram of the RESOLVE frequency domain electromagnetic system as currently implemented. The 9m long "bird" houses transmitter and receiver coils and is towed by the helicopter at a nominal altitude of 30 m above the ground (or obstacles such as trees). RESOLVE has 6 transmitter coils (horizontal co-planar coils, most sensitive to flat-lying features, range in frequency from 400 Hz to 140,000 Hz; there is one vertical co-axial coil with best sensitivity to vertical features, operated at 3,300 Hz). Magnetometers are placed on each side of the tail to measure the transverse magnetic gradient. Usually a magnetometer is also carried on the bird, as well as a laser altimeter to provide height above ground and an inclinometer to correct for changes in coil attitude.



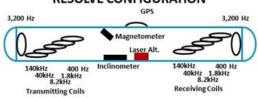


Figure 2 A picture of the RESOLVE system in flight (top) and a schematic diagram of the RESOLVE frequency domain EM system.

Method and Case Study - Area 1

The work flow is summarized by the following figure:



We provide an example of the workflow by reviewing the data and inversion results for one of the survey areas.

The Area 1 survey block contains 5 TSF, with strucutral reports available for three of these. A geology map of the area is shown in **Figure 3**. We will show examples from Block 1 of Area 1 (north east corner) where the geology consists of various metamorphic rocks.

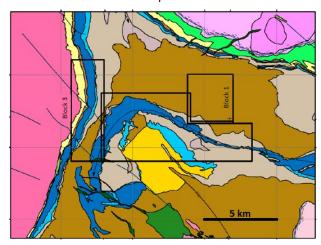


Figure 3 Geology map for the region with Block 1 being the top-right block (mainly metamorphic rocks, including shales, quartzites and metapelites).

RESOLVE FDEM and magnetic gradiometry data were acquired. In-phase and quadrature EM components were processed and preliminary interpretation products generated. These include apparent resistivity calculations and Differential Resistivity[™], a rapid conductivity depth algorithm.

For 3D inversion, cultural noise was first removed. Because of the tuned coils of RESOLVE, power line signal has a very different response in each of the frequency channels. **Error! Reference source not found.** shows the measured response in the vicinity of a strong power line. The highest frequencies (140 and 40 kHz) show very little response from the power line. The 8200 Hz shows some power line response, but the data is still very much interpretable. The lowest frequencies are swamped by power line noise, similar to what would be seen by a TDEM system.

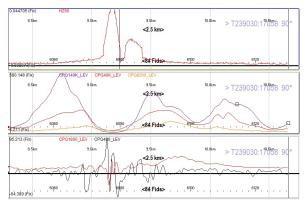


Figure 4 Power line monitor (top), 140, 40 and 8.2 kHz quadrature (middle) and lower freq. signal (1800, 400 Hz, bottom).

The 140 kHz apparent resistivity is shown in Figure 5. The tailing ponds are conductive in comparison with dams, in turn of intermediate conductivity when compared with the country rocks of higher resistivity. The apparent resistivity plot provide a check on the inversion process, identifying the main resistivity structures.

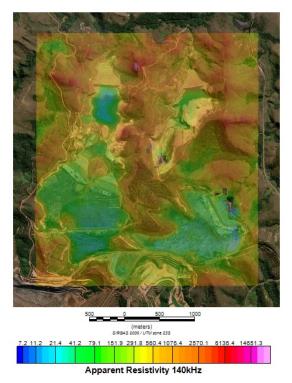


Figure 5 Apparent resistivity,140 kHz channel.

The survey was flown over a number of tailings ponds. In some instances – particularly older TSFs from other companies – it was not always possible to recover accurately recorded details about the TSF history. The geophysical surveys confirm existing records at many dam locations and in other areas provided very useful information about dam location and extent, and thickness of the tailings pond, as well as information about the surrounding geology.

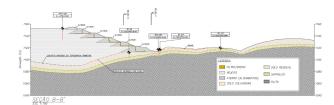


Figure 6 Structural section of dam #3. Note that the thickness of the tailings pond is inferred and not well defined.

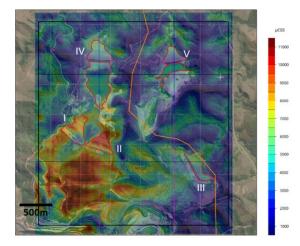


Figure 7 Magnetic susceptibility inversion; depth slice at 5 m below surface.

The magnetic inversion also clearly shows the tailings ponds and dams, and magnetic highs to the south east.

We can invert the RESOVLE in-phase and quadrature data with the magnetic data in a 3D joint inversion.

A number of tests were performed using both the magnetic and airborne EM data in direct joint and cooperative inversions. An important consideration in performing these joint inversions is the relative depth of penetration of the two techniques involved. The airborne EM data is expected to provide inversion results to approximately 100m depth. The magnetic data is sensitive the magnetically susceptible material from the surface to depths of many km's. In order to get meaningful results, the cooperative joint inversion approach was selected, where the magnetics is inverted using the EM resistivity result as structural constraint (Scholl et al., 2017).

An example of cooperative joint inversion is shown in Figure 8. The lower panel shows a slice through the unconstrained 3D magnetic inversion. The upper panel uses the EM structure in a joint cooperative inversion to better distribute the higher magnetic susceptibly material. As shown in Figure 7, high magnetic susceptibilities were correlated with the locations of some of the ponds, suggesting the pond fill contained minerals with magnetic susceptibility. The ponds were also more conductive than the surrounding geology. The cooperative joint inversion placed more susceptible material within the pond tailings, and the rest of the highly susceptible material was modeled as coming from the deeper geology.

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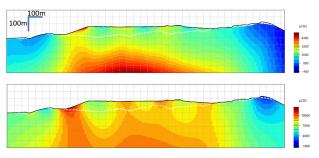


Figure 8 Comparison of the magnetic inversion using co-operative joint inversion with the FDEM data (top) and unconstrained magnetic inversion (bottom) for one of the dam sites.

Ground DC resistivity data was collected over a few of the dams in 2019, and inversion modeled before the current study. Figure 9 shows inverted resistivity for a series of ground DC lines over one of the dams. The tailings material appears as only slightly less resistive (~2,000 ohm-m) in absolute sense but still distinct from the highly resistive host rock (~10,000 ohm-m).

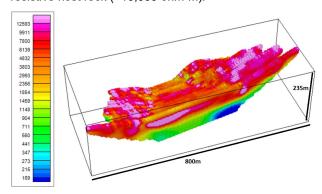


Figure 9 Legacy Inversion result from ground DC resistivity over dam #5.

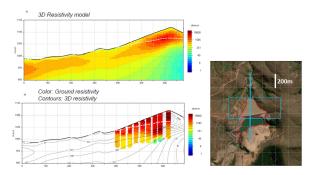


Figure 10 3D FDEM inverted resistivity (top). Below, DC resistivity (colour) and 3D inverted resistivity (contours). Section on map (right).

The DC resistivity lines were acquired across the dams, while the airborne lines were flown at approximately 90 degrees to this. Figure 10 provides a comparison the

inverted FDEM 3D resistivity with ground DC resistivity models conducted prior to this study. The airborne FDEM does not image the very resistive material at surface, likely because the skin depth in this resistor is very large so there is poor sensitivity. The FDEM shows a resistive wedge dipping to the north with approximately the same thickness as the DC resistivity. The FDEM inversion also shows a large conductor that is the tailings pond. Note that ground DC resistivity surveys, naturally, cannot be conducted over the pond itself.

Further Examples

Area 2

Area 2 contains numerous mines and several tailings ponds and dams. We focus on a dam that has DC resistivity lines over the rim, and in Figure 11 compare the DC inversion results with the 3D FDEM inversion. While the DC resistivity shows higher absolute values of resistivity, the trend shown within the dam is similar between the two methods.

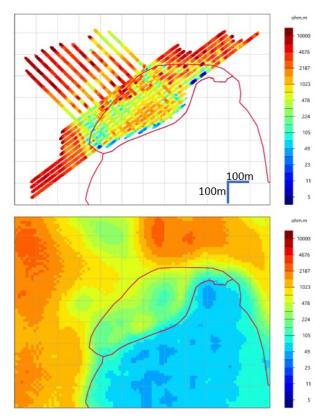


Figure 11 Comparison of ground DC resistivity (top) and 3D FDEM resistivity (bottom) at 20m below ground level.

Both show elevated resistivity values at the back of the dam at the front of the dam, curving away from the front face to the back of the dam towards the eastern extent. The central part of the dam shows a higher conductivity from both techniques.

Figure 12 shows a resistivity section through the dam and over the tailings pond. The tailings pond appears as a conductor, while the dam is a relative resistor. The basement here is very resistive. The 3D FDEM inversion is doing an excellent job mapping the thickness of the pond material.

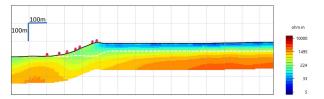


Figure 12 Section view of the Area 2 dam and pond (from Figure 13 and Figure 14). White line shows inferred pond depth from pre-dam construction SRTM data. Red symbols show locations of geotechnical drill holes.

Area 3

Figure 13 shows the magnetic data draped over a satellite image of Area 3. Here there is a large pond separated by an auxiliary dam. The pond shows a magnetic high, with significant variation. The dam at the southern end causes a distinct break in the magnetic response. Figure 14 shows the inverted resistivity at depths of 5 m, 20 m and 40 m. The tailings pond is conductive compared to the surrounding area. The FDEM inversion result is able to show variation in the thickness of the pond, which is deepest in the central part of the northern pond, exceeding 40 m. The dam itself is imaged as a resistor.

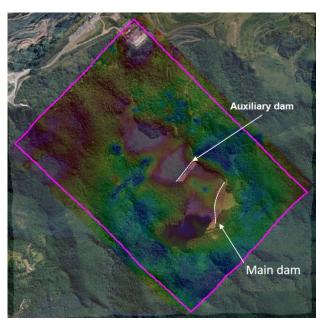


Figure 13 Magnetic response, Area 2 dams.

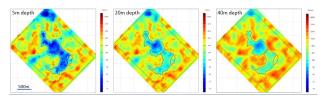


Figure 14 Inverted resistivity from FDEM data for depth slices at 5, 20 and 40m below surface.

Area 4

Here is an example of constraining the magnetic inversion with borehole data from a different survey area. Figure 15 shows the location of drill holes with F3GL information. The unconstrained magnetics inversion is compared to inversion constrained with borehole information in Figure 16. As the magnetic data has no inherent depth resolution, the borehole information is very useful in constraining the inversion process and adding detail to the inversion result.

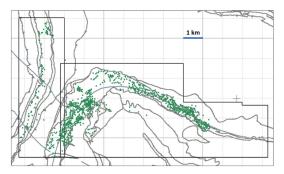


Figure 15 Location of drill holes with F3GL information.

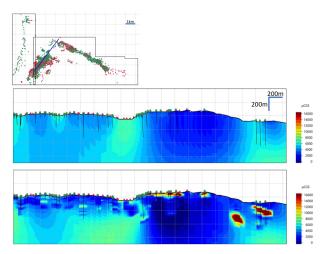


Figure 16 Profile location (top) with shallow (red) and deep (green) borehole locations, unconstrained (center) and constrained magnetic inversion result (bottom).

Conclusions

Vale is employing a pioneering, proactive strategy to manage historical mine sites. The results from a high resolution airborne frequency domain electromagnetic system and gradient magnetometer, inversion modeled in 3D, show that it's possible to efficiently image the subsurface, reliably. The resulting physical property models are being interpreted to better understand the geotechnical environment and history of dams and tailings ponds. The FDEM data in general agrees with ground DC resistivity but provides valuable advantages through spatially consistent and complete data coverage over the entire dams, tailings ponds and host rock system. The geophysical data has helped to confirm subsurface hypotheses at different locations and has provided additional information to further interrogate the existing geotechnical model.

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