Mechanism for an all around crust equatorial shear Carlos Eduardo de M. Fernandes - Universidade Federal do Rio de Janeiro - UFRJ

This paper was prepared for presentation at the 8th international Congress of the
Braellen Geophysical Society, held in Rio de Janeiro, Brazil, September 14-18,2003.

Contents of this paper were reviewed by the Technical Committ n of the f^o istional Congress of the Brasilian Geophysical Society and do not necessarily
ent any position of the SBGF, its officers or members. Electronic reproduction, or storage of any part of this paper for commercial purposes without the witten cond the Brazilian Geophysical Society is prohibited.

Abstract

Some geology and geophysics investigators have proposed, at the end of the 19th century (BIB. 1 and 2) disclosing evidences for it, that the Earth's crust had experienced, during the Mesozoic, an all around equatorial shear, that displaced the northern hemisphere westward relatively to the southern. During the 20th century, Prof. S. Carey (BIB. 3) and others (BIB. 4 and 5) reaffirmed the same proposition. However, no internal mechanism global-scale in range, has so far been advanced, to this author's knowledge, to justify it. This paper is an attempt, founded on the very effects of the Earth's rotation, respecting fundamentals, to prove that an equatorial shear stress, hence created, may account for the phenomenon, therefore providing an answer.

Introduction

Mechanical effects of Earth's rotation, as evaluated by the tensor field it generates at the near surface atmosphere and at the oceans are long known. Through extrapolations depthwise this tensor field may also be evaluated at the base of the crust and at any Mantle level.

The field comes, at each level, from the tensor component along a meridian plane projected down by the centrifugal repulsion and the likewise component along a parallel plane linked to the Coriolis backward drag.

Due to the huge masses that can deeply be mobilized. very high forces enter the picture and shear stress concentrations in the thin Crust naturally arise. Tensors moduli are computed, their resultant paths plotted and alobal-scale shear locations justified.

Parameters selection and evaluation

For an Earth shaped as an oblate ellipsoidal body, any meridian will have roughly for equation, the elipse

$$
\begin{bmatrix} R_{+} \cos \phi \\ 6378 \end{bmatrix}^2 + \begin{bmatrix} R_{+} \sin \phi \\ 6357 \end{bmatrix}^2 \approx 1
$$
 (1)

(in kilometers, where R_4 is the radius vector and ϕ the latitude) and a parallel at latitude ϕ , the circumference

$$
x^2_{\phi} + y^2_{\phi} \approx R^2_{\phi} \text{.} \cos^2 \phi \tag{2}
$$

The slow Earth's rotation, $w \approx 73*10^{-6}$ rad/s, projects at its surface (h=0) or at any depth (h>0), the following pair of tensors: one, tangent to a meridian line, confined to the respective radial plane, directed from the poles to the equator ($M_{0,4}$, $M_{n,4}$) and the other, tangent to a parallel line, confined to the respective orbital plane, directed retrogradely to the planet's rotation, i.e. westwards, (Po.,, $P_{h,4}$).

Figure 1 : Mo., and Mk, loci in the Meridian plane

Figure 2 : Ps, and Ps, loci in two parallel planes

These tensors have for expression, respectively, $M_{0, \phi}$ = $=w^2$ R₄. cos₄. sen₄ = 1/₂ w² R₄. sen2 ϕ $M_{h,4} = 1/2$ w². (R₄-h). sen2 ϕ (3) as tangencial components of the respective centrifugal distinguished component
repulsion, and
 $P_{0,4} = 2w^2$. R_+ . cost
 $P_{h,4} = 2w^2$. $(R_+ - h)$. cost (4)

in accordance with the original Coriolis deduction.

(at the S'ern hemisphere)

They yield, then, at each point the resultant tensor $T_{0, \phi}$ whose construction is shown (vector-like, since M_{0,4} and $P_{0, \phi}$ can only act on a common area, therefore in a shear fashion: see proof).

 (M) (M)

(at the N'ern hemisphere)

(at the equator) Figure 3

Proof: if $M_{0,4}$ and $P_{0,4}$ where normal to their respective infinitesimal areas, then

 $T_{0,+}$ 13 = M₀, 12 sen θ + P_{0, +} 23 cos θ , or

$$
T_{0, \phi} = M_{0, \phi}
$$
 12/13. sen θ + P_{0, \phi} 23/13. cos θ =

= $(M_{0,4}+P_{0,4})$. sen 20. It is known that at ϕ =

=0, $M_{0,0}=0$ and $P_{0,0}=T_{0,0}$; in the above expression this would lead to sen20=2, which is impossible.

The following Table of Results summarizes, at 15° steps in latitude, tensors moduli and azimuths, at the surface and at 30, 700 and 3000 km in depth. The unit system is the CGS for tensors: they are in gats (or dynes/gram), instead of cm/s², to emphasize the fact that no free movements take place.

Table of Resuts

Eighth International Congress of the Brazilian Geophysical Society

 $\overline{2}$

Remarks: in this table, $\frac{1}{T_{n,4} = M_{n,4} + P_{n,4}}$, or
 $T_{n,4} = (M_{n,4}^2 + P_{n,4}^2)^{1/2}$, is the resultant tensor and
 $\theta_{n,4} = \tan^{-1} (P_{n,4} M_{n,4})$; being $\theta_{n,80^2} = \tan^{-1} 2 = 63.4^{\circ}$, its azimuth.

Interpretation of Results-Final Remarks

On the Geoid's surface (coordinates $0,\phi$) T_{0,4} can be represented by the family of lines shown below (drawn unavoidably distorted due to the map's projection). At each line the azimuths (θ) vary from 63.4° (at ϕ = 90°) to 90° (at $\phi = 0$ °) and the tensors moduli, respectively, from

zero (0) to 6.75 gal; that pattern ought to repeat itself, undistorted, on the spheroidal surfaces below at 30,700 and 3000 km, since $T_{n, \phi}$ is, with due approximation, a lineal function of depth (h). At the near surface atmosphere and at the oceans being both fluid, free within limits to move, they do so accordingly to an imprint of $T_{0,*}$ on a broad
scale. Deviations and discrepancies arise, due to temperature anomalies and physiographic irregularities both above and below sea level. This known correlation is, at this point, recalled to fundament what follows.

Figure 5

The map below is a reproduction of a set of major shears proposed by several authors The name "megashear" was
used originally by S.Carey (BIB 3) to

designate "a strike-slip fault whose horizontal displacement exagginate a suite with the thickness of the crust" (sic); the
largest would be the "Tethyan Shear System", near
equatorially located and all around the Globe:

Figure 6

We concur with the assumption that continents are welded to their own Mantle: continental drifts are Mantle drifts. Tensors $T_{30,4}$, $T_{700,4}$ or $T_{3000,4}$ line patterns are those of $T_{0,4}$, and they act drift-like, not push-like. It becomes, than, striking the fact that the most extensive shear be equatorial located, where $T_{h,0}$'s are maxima: being tensors, The immediately at the northern side of the

equator does not interfere either construtively-or not with The at the southern surroundings; they should coexist. However, if the intervening masses differ, as apparently they do (see figure), the ensuing forces will also differ, and hence a shear stress, on a near 90° dip surface across the thin Crust, trending equatorially, would prevail. This fits the megashear geometry and justify it.

The magnitude of the associated stress, equatorially aligned, can be evaluated from the following fundamental relations.

Driving forces (f_N, f_s) at each hemisphere, \blacktriangleright projectd at the equator,

 $f_N \approx 2\pi$ * \overline{R}^2 * t_N * δ * $\overline{I}_{0,*}$ * $\cos 70^\circ$

 $f_s \approx 2\pi * \bar{R}^2 * t_s * \delta * \bar{T}_{0,*} * \cos 70^\circ$

where R is an average radius ≈ 6367 km $\approx 6367x$ 10⁵ cm; t_w and ts, average crust thicknesses at each hemisphere;

 (5)

 δ , estimated crust mean density ≈ 2.8 g/cm³;

 $\overline{T_{0,\bullet}}$, average value for the resultant tensor at each hemisphere ≈ 3.5 cm/s²;

70°, approximate mean angle between $\overline{T_{0,4}}$ and the equator line.

These forces act along the same westward direction at the equator, but being different in value (since $t_N > t_S$), the force differential $\Delta f = f_N - f_S$, divided by the common equatorial contact area(assumed with 90°dip), S≈ 2x*Ro*ts will yield the acting differential shear,

$$
\triangleright \tau_a \approx (\Delta f/S) \approx 2\pi * (6367 * 10^5)^2 * 2.8 * 3.5 \times 0.342 * (t_N - t_S) \approx 2.13 * ((t_N - t_S) - 1) * 10^9 \text{ dynes/cm}^2
$$
 (6)

Now, from the geophysical literature (BiB.5) the average shear resistance of the rock-forming crust at depth may be taken between the limits

0.14 * 10⁹ dynes/cm² ≤ τ ≤ 0.42 * 10⁹ dynes/cm² and by assigning any of these possible ratios

$$
t_w t_s = 1.10 \qquad \tau_a \approx 0.21*10^3 \text{ dynes/cm}
$$

 $\tau_a \approx 0.32*10^9$ dynes/cm² $t_N / t_S = 1.15$

 $\tau \approx 0.43*10^9$ dynes/cm². twks = 1.20

a favourable numerical credibility is achieved in support of this paper's thesis. Finally, it should be remarked that the equator line, as can be deduced by inspection of Ta_{ϕ} geometrical pattern, represents a hiatus in the N-S, or meridian transmission of T_{0,0}: this would indicate that the N'em hemisphere does not push against the S'em, or viceversa, independently of any rock continuity they may possess.

Bibliography

- 1 GREEN, W.L,1887: The Earth's Features and Volcanic Phenomena. Hawaian gazette,
- 2. Honolulu. PRlNZ, W.,1899: L'échele Reduite des Expériences Géologiques. Soc. Belge
Astronomique, Bulletin, Belgic.
- s. Astronomique, Bulletin, Belgic. CAREY, S.W.,1976: The Expanding Earth. Elsevier Sc. Pub.Co. Amsterdam, HoIland.
- 4. HIDE, R., 1969: Interaction between the Earth's Liquid Core and Solid Mantle. Nature,
- 222;1055-6
5. JEFFREYS, H., 1967:Inelastic Processes in the Shel. Roy. Astr, Soc., GeophYS. Jour. 14; 1-4
- SYNGE, J. L E GRlFFITH, BA., 1949. 6. Principies of Mechanics. Mc Graw-Hill. NY,US. JAGER, C., 1969: Elasticity, Fracture and Flow. Methuen and Co. Ltd. London. England.

Acknowledgments

To Huaila Fonseca Ayres a geology student at our University (U.F.R.J.) and Lucia Nunes P. Barsch, the Engineering Geology Setor's secretary, for ali the help to edit this manuscript.