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STUDY OF ELASTIC PROPERTIES AS FUNCTION OF TEMPERATURE IN ANISOTROPIC CRACKED MEDIA: AN ULTRASONIC APPROACH

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ABSTRACT. The study of seismic wave velocities and their related anisotropy parameters provides important tools to the study of Earth's subsurface. The analysis of temperature's influence in the elastic properties of synthetic rocks, e.g., may be a good alternative for modelling the response of elastic properties of rocks surrounding deep wells and/or that are submitted to the process of steam injection. In order to understand the effects of temperature on the elastic parameters, in this work we constructed four porous synthetic sandstones, one representing an uncracked medium and three cracked samples. They were submitted to variable temperatures, while their elastic properties were calculated during heating and cooling processes. It was noted that P- and S-wave velocities decreased with increasing temperature. This behavior may be caused by changes in the background stiffness and by the generation of secondary cracks inside the samples. However, P- and S- anisotropy parameters were not affected by the changes in temperature. This can be an indicative that randomly distributed secondary cracks were created. Other indicative of secondary cracking formation can be associated to the difference between seismic velocities during the heating and cooling processes.

Keywords: cracked media, anisotropy, temperature dependence, elastic parameters.

RESUMO. O estudo das ondas sísmicas e de seus parâmetros anisotrópicos associados fornece ferramentas importantes para o estudo da subsuperfície da Terra. A análise da influência da temperatura nos parâmetros elásticos de rochas sintéticas, por exemplo, pode ser uma boa alternativa para modelar a resposta das propriedades elásticas de rochas circundando poços profundos e/ou que são submetidas a processos de injeção de vapor. Para entender os efeitos da temperatura nos parâmetros elásticos, neste trabalho construímos quatro arenitos sintéticos porosos, um representando um meio não-fissurado e três amostras fissuradas. As amostras foram submetidas a temperaturas variadas, enquanto suas propriedades elásticas foram calculadas ao longo dos processos de aquecimento e resfriamento. Notou-se que as velocidades das ondas P e S diminuíram com o aumento da temperatura. Esse comportamento talvez seja causado por variações na rigidez do *background* e pela formação de fissuras secundárias nas amostras. Entretanto, os parâmetros anisotrópicos P e S não foram afetados pelas mudanças de temperatura. Isso pode ser considerado um indicativo de que fissuras secundárias distribuídas aleatoriamente foram formadas. Outro indicador da formação de fissuras secundárias pode ser associado à diferença entre as velocidades sísmicas durante os processos de aquecimento e resfriamento.

Palavras-chave: meios fissurados, aniostropia, dependência de temperatura, parâmetros elásticos.

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INTRODUCTION

Seismic wave velocities are known to be dependent on many factors, such as overburden and pore fluid pressure, lithology, porosity, composition and amount of pore fluids, pressure and temperature. Among these, lithology, porosity and pressure are considered to be the primary variables in affecting velocities (Timur, 1977). Nonetheless, the temperature effect plays a major role on seismic velocities under conditions of high temperatures, such as those found either in the steam injection process or in deep wells. Another factor of importance to seismic wave velocities is anisotropy. Vertical transversely isotropic (VTI) medium have been the subject of many studies along the years due to its great occurrence in sedimentary rocks and its importance to the oil industry. Thomsen (1986) has shown that, in the case of weak anisotropy (10%-20%), the anisotropy of a medium can be described using three parameters: ε, γ and δ.

Investigating the temperature and pressure effects on seismic velocities of halite salt, Yan et al. (2016) concluded that the temperature effect on the velocities is dominant relative to the stress effect. Timur (1977) measured acoustic velocities in sandstone and carbonate samples as a function of temperature and found that the average decrease for a 100°C rise in temperature was 1.7% for compressional wave velocities and 0.9% for shear wave velocities. Kern (1978) investigated the effect of high temperature and high confining pressure on compressional wave velocities in rocks of different types and found that the velocities generally decrease with temperature, but the rate varies according to the rock type. Furthermore, he concluded that, at higher isobaric pressures, the velocity drop as function of temperature decreases and the velocity-temperature relationships tend to become more and more linear functions.

Other studies have observed a decrease in elastic modulus (Zhang et al., 2009; Brotóns et al., 2013) with increasing temperature, which accounts for a decrease in seismic velocities. Thermal gradient cracking has been found to be the dominant mechanism for fracture formation in experiments conducted at low confining pressure, resulting on the decreasing of seismic wave velocities with temperature (Jansen et al., 1993). Most of the previously mentioned works did not study the anisotropy dependence on temperature. Kern & Fakhimi (1975), however, studied the influence of temperature on anisotropy in various metamorphic rocks under high pressure conditions, and observed that fabric-induced seismic anisotropy is generally unaffected by temperature. On this study, however, Kern & Fakhimi (1975) analyzed only P-wave anisotropy.

This study aims to further investigate the temperature influence on compressional and shear waves traveling in a VTI synthetic media, as well as the dependence of the Thomsen's parameters (ε, γ and δ) with respect to temperature. Measurements of traveltime were performed, through the ultrasonic method under atmospheric pressure, as a function of temperature (ranging from 25°C to 130°C) in four cracked samples constructed using the technique developed by Santos et al. (2017). Each synthetic sample has a different crack density, which characterizes different degrees of anisotropy, allowing us to analyze velocity changes with respect to temperature and crack density.

METHODOLOGY

The construction of the synthetic rock samples as well as the ultrasonic measurements were performed at the Laboratory of Petrophysics and Rock Physics – Dr. Om Prakash Verma (LPRP) at the Universidade Federal do Pará, Brazil. The following methodology description corresponds to the production of fractured or cracked synthetic sandstones.

Sample preparation

We employed the methodology developed by Santos et al. (2017) for crafting synthetic VTI samples, which is based on: cement, sand, styrofoam and paint thinner. Four samples were constructed with different crack densities. Table 1 shows the number of cracks per layer and the crack density of each sample. In order to ensure that the samples had the same composition, so that the differences in elastic properties were caused only by different crack densities, all samples were crafted simultaneously at the same mold, as can be seen in Figure 1.

Table 1 – Crack densities of the samples evaluated using Eq. 1.

Sample	Cracks per layer	Crack Density
REF	–	–
L1	14	4.0%
L2	21	6.0%
L3	28	8.0%

The mortar used in the samples construction was composed of 70% sand and 30% cement. Moreover, the proportion of water was kept constant to ensure the samples homogeneity in composition. The cracks were disposed on layers constructed one
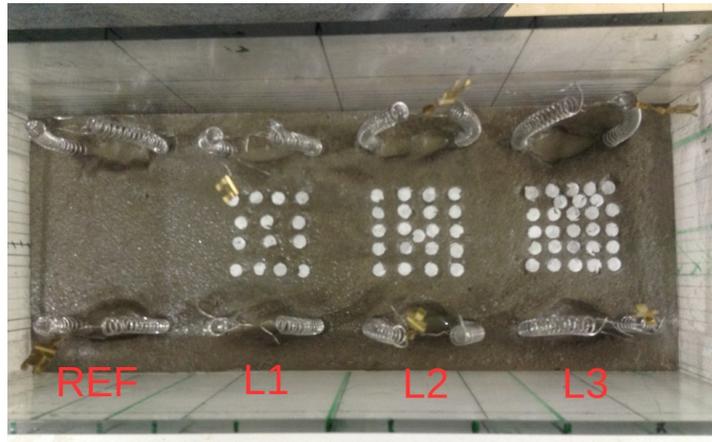
Figure 1 – Crafting process of the samples showing the styrofoam discs disposed at the top of the mortar layer and the resistances positioned inside of samples. The Figure 2 shows all samples cut and after the chemical launching to remove the styrofoam discs.
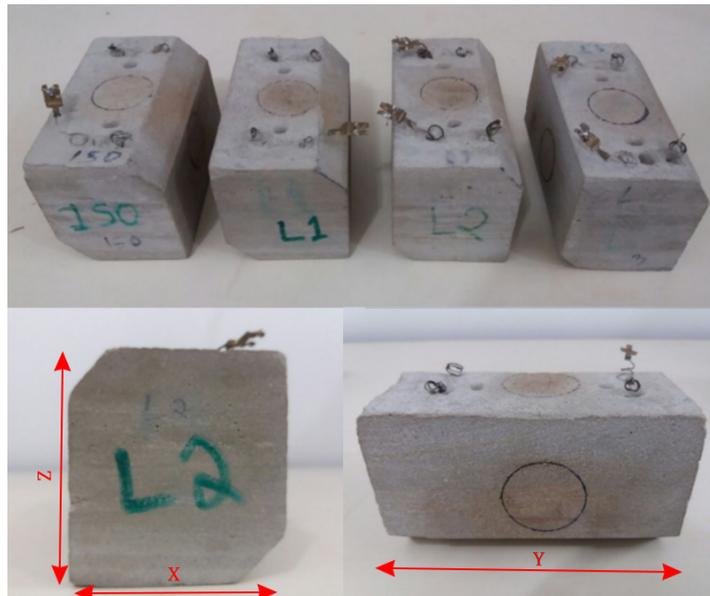

Figure 2 – Image of the four samples constructed and analyzed in this work.

at a time. At first, we filled the mold with a 2 cm layer of mortar and then placed penny-shaped styrofoam sections that were 0.5 cm in diameter and 0.1 cm thick (Fig. 1). We then filled another 0.5 cm layer of mortar and placed another set of penny-shaped cuts at the same area (4 cm x 4 cm) as before. This process was repeated until we had 7 layers with cracks spaced by 0.5 cm. Finally, we filled the last 2 cm thick layer totalizing a height of 7 cm.

Right after filling the first layer, nichrome resistance pairs were placed in the regions of each sample (Fig. 1). Those resistances are part of the temperature control system which is described in Silva Sobrinho et al. (2017). To enable for the 45° measurements, the location where the resistances were placed at the mold was chosen taking into account the 45° cuts made afterwards on the edges of the samples. As a matter of fact, the dimensions of the samples were chosen so that a cross section of the XZ-plane would be approximately a square, which in turn would enable propagation in the 45° direction.

The curing time of mortars produced with cement was about eight days (Garcia et al., 2011). Once the curing time was over we had the samples sawed apart. From Santos et al. (2017), to create

penny-shaped voids, a chemical leaching using paint thinner was performed. The samples were left immersed in paint thinner for around 24 hours and some samples, after the leaching process, were cut in order to verify the absence of styrofoam at the cracks. As the styrofoam is a polymer made by 98% of air and 2% of polystyrene (Poletto et al., 2014), after the leaching a practically zero amount of polystyrene mass remains inside the fracture. The samples crack density was estimated by the modified Hudson (1981) 's equation, given by

$$\varepsilon = N_i \frac{\pi h_i r_i^2}{V_m},\qquad(1)$$

where N_i is the total number of penny-shaped inclusions, h_i is the aperture (thickness) of inclusion, r_i is the radius of the inclusion and V_m is the model volume (only) occupied by cracks.

Ultrasonic setup and Data acquisition

The ultrasonic measurements were performed using the LPRP ultrasonic system with the same pulse transmission technique as performed in De Figueiredo et al. (2013), Santos et al. (2016) and De Figueiredo et al. (2018). The sample rate by channel for all measures of P- and S-waveforms was 0.005μs. Figure 3 shows the ultrasonic system. The system is compound by a pulser-receiver 5072PR and a preamplifier 5660B by Olympus, a 50 MHz USB oscilloscope by Handscoope, and two transducers of 1 MHz (P-wave) and 500 kHz (S-wave) by Olympus.
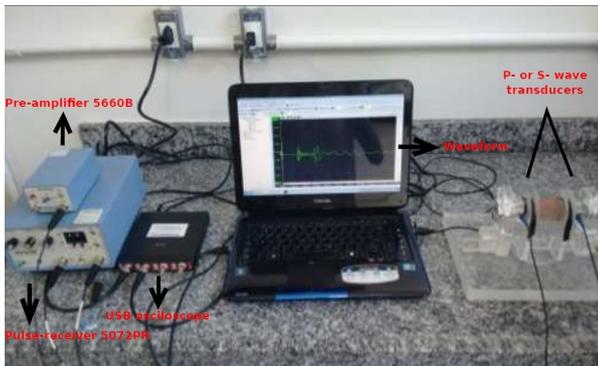

Figure 3 – Ultrasonic measurement system showing the pre-amplifier, the oscilloscope, the pulse-receiver and the P-and S-wave transducers.

The measurements were realized in the directions 0°, 45° and 90°, being 90° the direction parallel to the cracks, 0° the direction perpendicular to the cracks and 45° diagonal to the cracks. Figure 4 depicts the directions of propagation and polarization assigned on this study. All the measurements were performed with the sample temperature varying from from 25°C to 130°C. Temperature step was 15°C, which resulted in a total of

8 temperature values. In the direction 90°, two polarizations (z and y) were used in order to evaluate V_{s1} and V_{s2}. The measurements were carried out during the heating and the cooling for each sample, in an attempt to investigate the hysteresis of elastic wave velocities caused by temperature change. To guarantee a good contact between transducer and samples, we used honey as type of couplant.
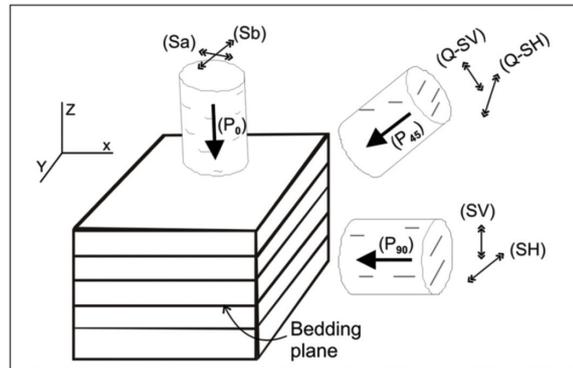

Figure 4 – As shown at Figure 2, our samples has a rectangular parallelepiped format. However, here we use a illustration of VTI medium to better show the sketch of the wave propagation direction and polarization to obtain the P- and S-waveform records. This image was modified from Martínes & Schmitt (2011).

At the beginning of the heating process the power was kept low (2-4%) to avoid the overshooting of the goal temperature. The power was increased with the increasing sample temperature. At the maximum temperature, the power was about 10-11%. Afterwards, the power was gradually decreased in order to take ultrasonic measurements during the cooling process. Figure 5 depicts the experiment during the data acquisition.
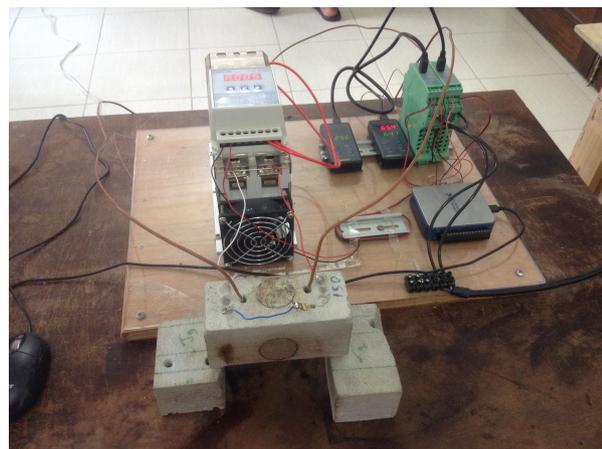

Figure 5 – Temperature control system acting on REF sample (isotropic medium) during data acquisition.
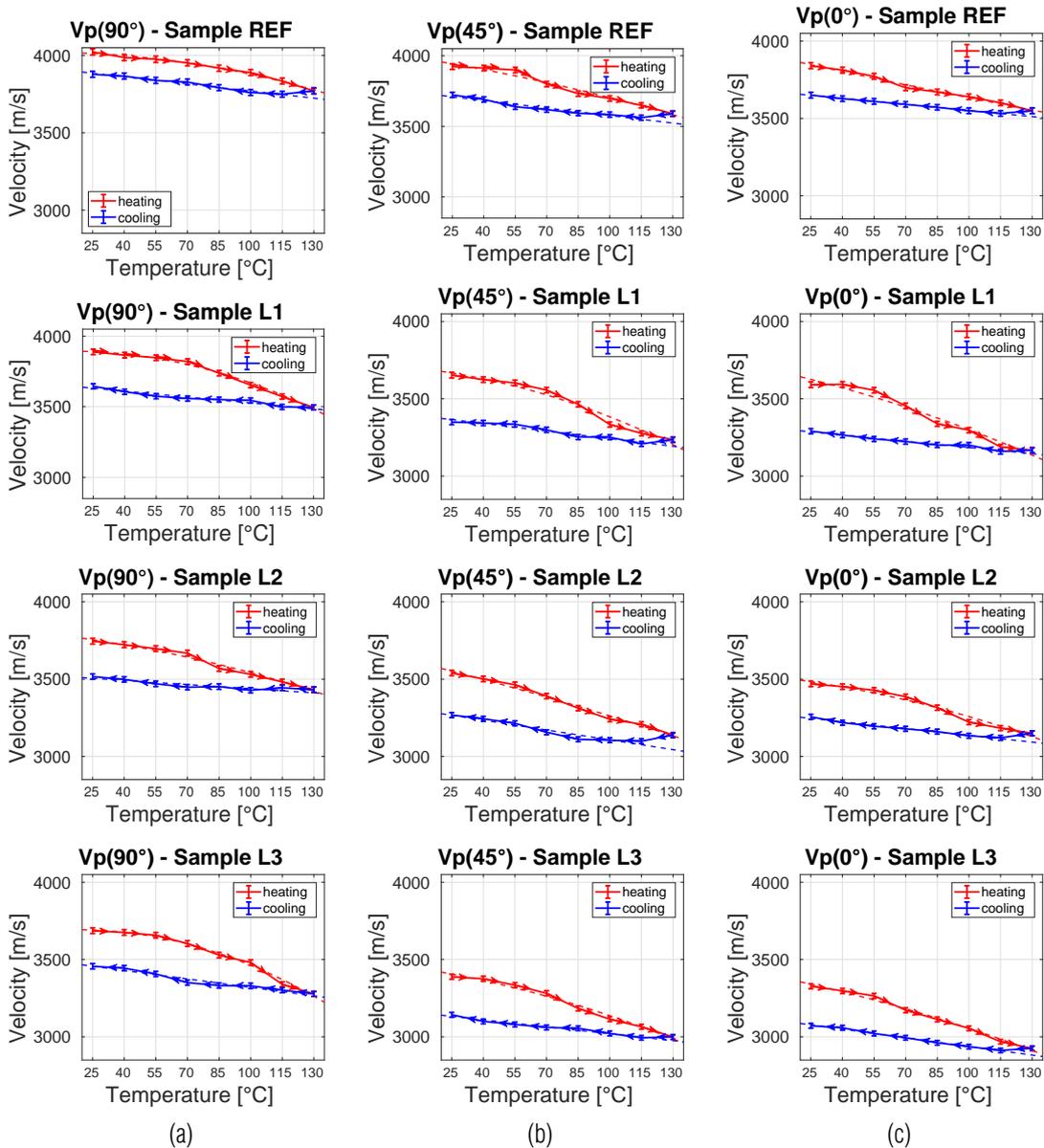
Figure 6 – P-wave velocities ((a) Vp_{90}, (b) Vp_{45} and (c) Vp_0) as function of temperature for samples REF, L1, L2 and L3 for heating and cooling cycles.

To estimate P-wave velocities, we used the relation given as:

$$V_P(\theta) = \frac{L_P}{t_P(\theta) - \Delta t_{delay}}, \qquad (2)$$

where L_P is the distance of P-wave propagation, $t_P(\theta)$ is the transmission time as a function of the angle with respect to the Z-axis of a P-wave, and Δt_{delay} is the delay time due to the S-wave transducers.

For S-wave velocities, the equation is similar to P-wave. They are given by

$$V_S(\phi) = \frac{L_S}{t_S(\phi) - \Delta t_{delay}}, \qquad (3)$$

where L_S is the distance of S-wave propagation, $t_S(\phi)$ is the transmission travel time as a function of the angle polarization.

For experimental velocities estimation, it was considered errors due to measurements of length (margin of error = $\pm$ 0.02 cm) and time travel picking (margin of error = $\pm$ 0.02 microsecond (μs)). The uncertainties for P- and S-wave
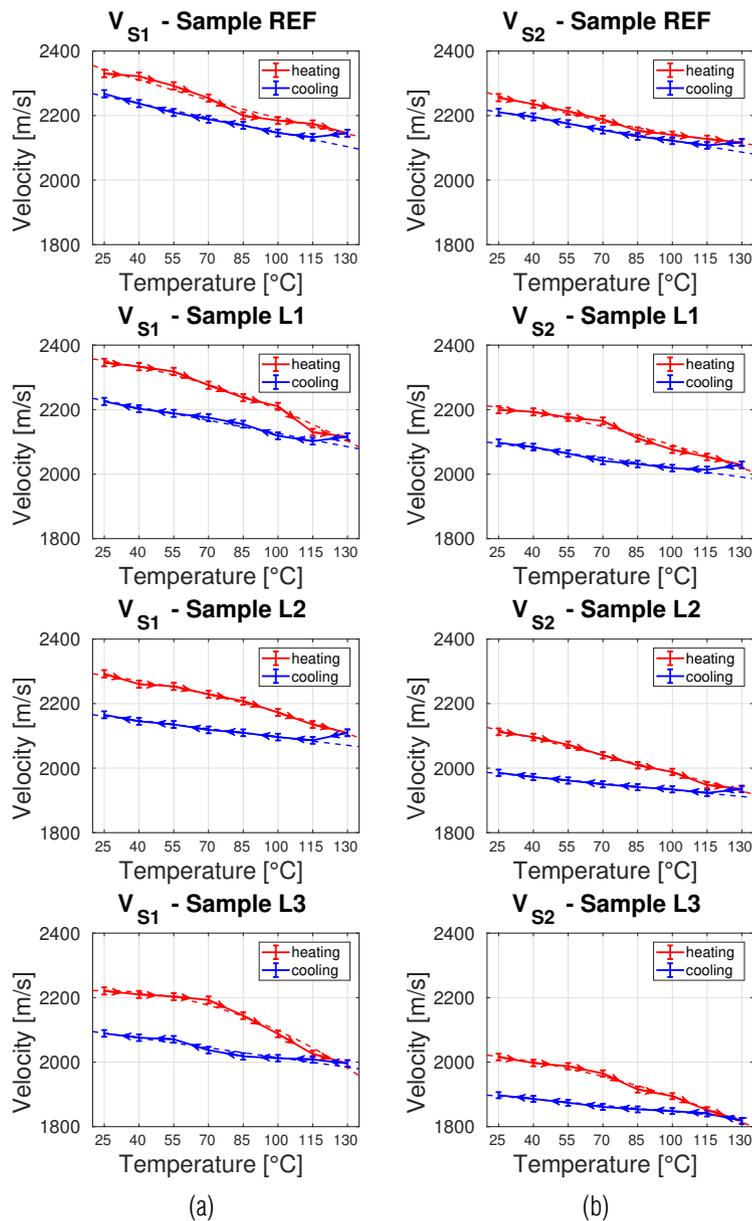
Figure 7 – S-wave velocities ((a) V_{s1} and (b) V_{s2}) as function of temperature for samples REF, L1, L2 and L3 for heating and cooling cycles.

velocities estimate ranged from $\pm$ 20 to $\pm$ 35 m/s and $\pm$ 18 to $\pm$ 28 m/s. Related to the temperature, the margin of error was $\pm$ 5°C for both heating and cooling process.

RESULTS

Figures 6 and 7 show, respectively, the variation of P- and S-wave velocities as a function of temperature, for all propagation and polarization directions. They are related to all the samples in study and show the behavior of compressional and shear velocities during both the heating (red curves) and cooling (blue curves) of

samples. For all the cases it is observed a decrease in velocities with the increasing temperature. Also, it is evident the existing of hysteresis in all the samples for all velocities, i.e., there is a variation of velocity for the same temperature when the sample is heating and cooling. Using the support of previous works to interpret these behaviors, such as the works of Timur (1977), Kern (1978) and Jansen et al. (1993), it can be said that they may be caused by the generation of microcracks inside the samples (secondary cracking), the coalescence of those microcracks and the decrease of the background stiffness of the samples. As

Figure 8 – Sample L2 before and after heating. It can be seen the presence of fissures on the surface caused by the coalescence of microcrasks.
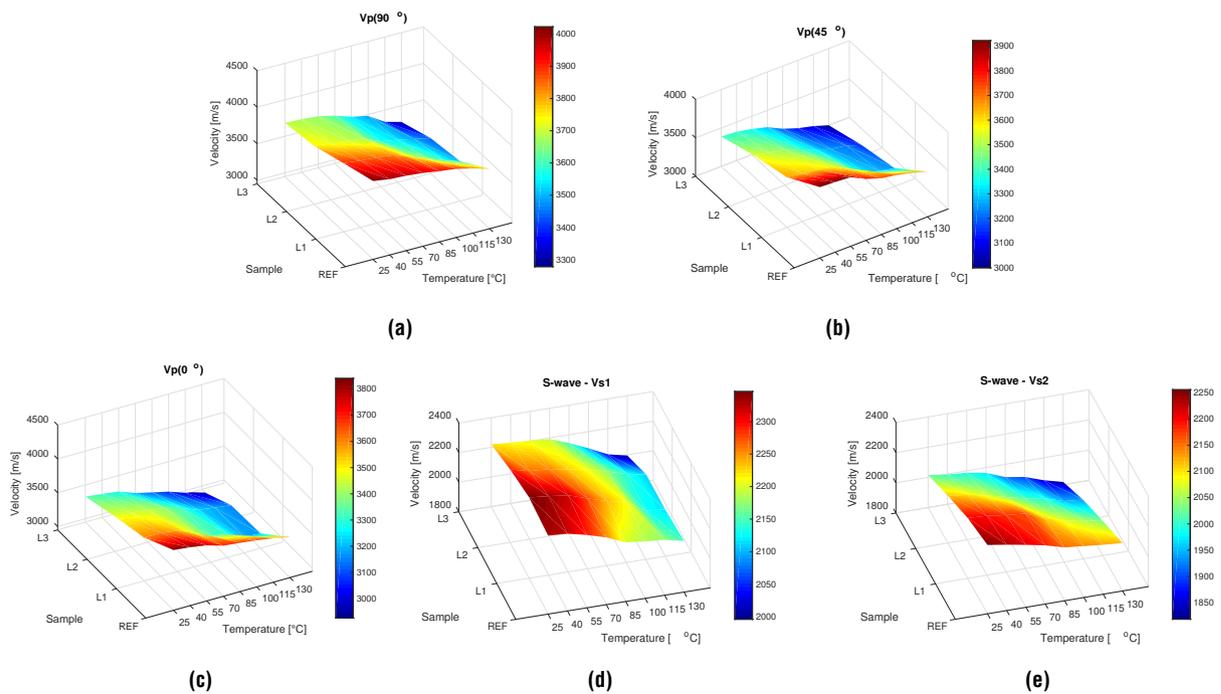

(a)

(b)

(c)

(d)

(e)

Figure 9 – Elastic velocities ((a,b,and c) P and (d and e) S) as function of temperature and crack densities shown by maps graphics. For all maps the highest velocity occurs when the temperature is 25° and the crack density is 0. The colorbar unity is m/s.

expected, the opposite behavior occurs when the samples cooled. Since the microcracks generated by the heating process do not disappear when the samples are cooled down, it can be inferred that the behavior showed by the blue lines, where the velocities increased with the decreasing temperature, are solely or mostly related to the variation of background stiffness. Another indicatives of the secondary cracking and cracking coalescence is the fact that the velocities never return to their initial values after the cooling process and the increasing in porosity for all the

samples after the experiment, as it can be seen in Table 2. The coalescence of microcracks created fissures on the outer edge of the samples, as it can be seen in Figure 8 that shows sample L2 before and after heating. On average, P-wave velocity decreased 11% for a 105°C rise in temperature, whereas S-wave decreased on average 9% for the same temperature gradient.

Figure 9 shows the variation of all the velocities as function of both temperature and the crack density, representing by the sample labels. Figure 9 concatenates all velocity data from
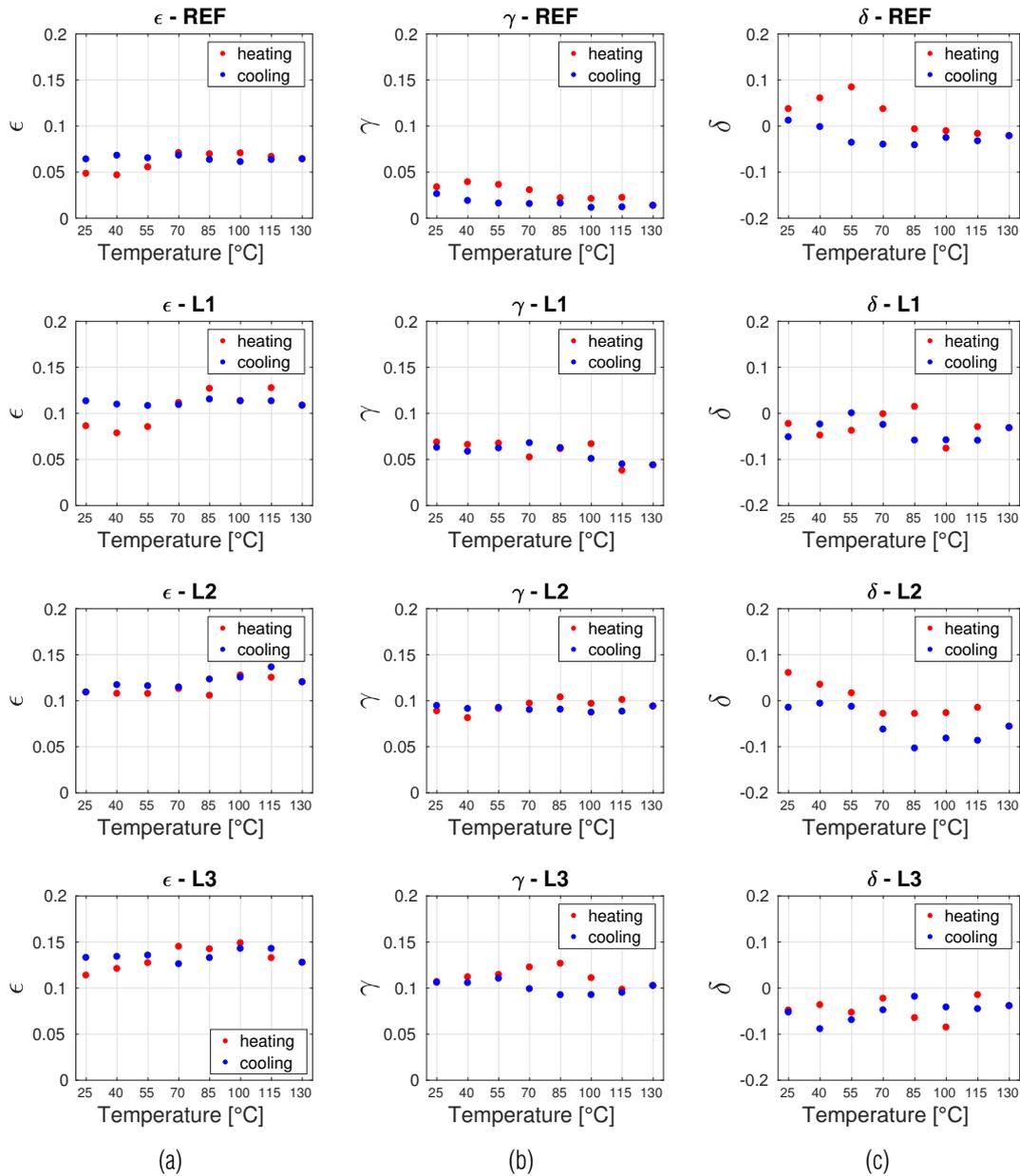
Figure 10 – Thomsen's ((a) ε, (b) γ and (c) δ) parameters as function of temperature for samples REF, L1, L2 and L3 for heating and cooling cycles.

Figures 6 and 7 during the heating process. It can be seen that the gradient of velocity as a function of temperature and crack density points in the direction of minimum temperature and crack density, which is a expected behavior since velocity decreases with increasing temperature and increasing crack density.

Table 2 – Initial porosity, final porosity and secondary crack porosity (after heating and cooling) of all samples used in this study.

Sample	Initial Porosity	Final Porosity	Induced Porosity
REF	8.2%	18.0%	9.8%
L1	12.0%	21.0%	9.0%
L2	14.0%	23.5%	9.5%
L3	15.5%	25.5%	10.0%

Figure 10 shows the variation of the Thomsen's parameters (γ, ε and δ) as a function of temperature, for sample L2, during both the heating and cooling processes. The γ and ε values increase as the number of cracks of the samples increase, which is an expected behavior. In most of the cases, the γ and ε values show very close heating and cooling values, which is consisting with the observations made by Kern & Fakhimi (1975) that temperature does not noticeably reduce fabric-induced anisotropy. This may be caused by the fact that P-wave velocities in the $0°$ direction decrease at approximately the same rate as the $90°$ direction; the same occurs for S-wave velocities for different polarizations. This is an indicative that the secondary cracks generated during the heating process are randomly oriented. As for the δ parameter, it does not seem to show any correlation with temperature variation.

CONCLUSIONS

An ultrasonic procedure was performed on synthetic samples representing isotropic and cracked sandstones. They were constructed using cheap materials and a very simple construction methodology. The waveforms were acquired from the samples for different temperatures, which was possible by using a temperature controlling system. From the analyzes of the results, the following conclusions were obtained:

- The increasing temperature decreases the background stiffness and generates secondary cracks inside the samples.

- P-wave velocities had on average a 11% decrease for a $105°$C rise in temperature, and P-wave velocity decreased approximately at the same rate on all directions of propagation.

- S-wave velocities had on average a 9% decrease for a $105°$C rise in temperature, decaying at the same rate for both polarizations.

- Anisotropy parameters γ and ε presented very close values both for heating and cooling. They do not show a satisfactory pattern with the increasing and decreasing temperature, making their use as tools to model elastic properties as a function of temperature not applicable.

In addition, it is worth to highlight the importance of this work as the first approach using porous synthetic rocks in the analysis of the influence of temperature in the behavior of elastic properties. The same approach may be repeated using rocks with different mineralogical composition and/or anisotropy symmetries.

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